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# NRCS CURVE NUMBER CALIBRATION USING USGS REGRESSION EQUATIONS

by

Charlotte Mecham Adsero

A thesis submitted to the faculty of

Brigham Young University

in partial fulfillment of the requirements for the degree of

Master of Science

Department of Civil and Environmental Engineering

Brigham Young University

August 2008

## **BRIGHAM YOUNG UNIVERSITY**

## GRADUATE COMMITTEE APPROVAL

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This thesis has been read by each member of the following graduate committee and by majority vote has been found to be satisfactory.

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and Technology

#### **ABSTRACT**

## NRCS CURVE NUMBER CALIBRATION USING USGS REGRESSION EQUATIONS

### Charlotte Mecham Adsero

## Department of Civil and Environmental Engineering

### Master of Science

The Curve Number (CN) method of estimating the direct runoff response to rainfall events was originally developed in the 1950's primarily for agricultural purposes in the mid-western United States. The accuracy of the CN method is greatly affected by variation of the soil type and land use of the region. Curve Numbers developed for a given region are not appropriate for application in other regions. In order to produce reliable, consistent results, Curve Numbers must be calibrated for the area in which the CN method is to be applied.

Calibration is ideally accomplished by direct measurement using several rain and stream gauges within a watershed. Gauged data, however, is not always available or easily obtained. A more feasible method of calibration is therefore necessary for broad application of the CN method.

The purpose of this study is to develop a method of CN calibration that can be easily applied to regions where no gauged data is available using the United States Geological Survey (USGS) regression equations. In this study, the peak flow values estimated using the regression equations were used in conjunction with a dimensionless hydrograph to compute runoff volume. The National Oceanic and Atmospheric Administration (NOAA) rainfall grids were used to estimate precipitation. Given the rainfall and runoff, a Curve Number can then calibrated through back-calculation.

The method of CN calibration using the USGS regression equations was applied to nearly 60 watersheds in the state of Utah for this research. The calibration results obtained using the regression equations were compared to other CN calibrations developed using gauged data. Calibrations performed through the use of the regression equations were quite consistent with calibrations obtained using measured data. To ensure the validity of the application of this method in other regions, more comparisons to results obtained using measured data should be further pursued.

## **ACKNOWLEDGMENTS**

I wish to thank Dr. E. James Nelson, my advisor for all his time, expertise and guidance. Thanks to Dr. Pete Hawkins, Dr. Rollin H. Hotchkiss, and Dr. A. Woodruff Miller for their assistance and review of my work, and the Utah Department of Transportation for their support. To my husband Matt Adsero, and my family; thank you for all your love and support.

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## 1 Introduction

An accurate understanding of surface hydrology is a critical element in the design of roads, bridges, culverts, and other structures. The ability to predict and design for future rainfall and surface runoff volumes is not only helpful in ensuring adequate drainage capacity of these structures, but also economically important. Predictions of runoff volume that are too high could result in the over-design of the structure and thus, unnecessary expenditures. Predictions that are too low could result in catastrophic damages to the structure and costly repairs.

Several methods for estimating the rainfall-runoff relationship have been developed to assist in the appropriate design of these types of structures. One of the more common and widely used methods is the Curve Number (CN) method (Hawkins *et al.*, 2006). The CN method was developed in the mid-1950's by the Soil Conservation Service (SCS) now known as the Natural Resources Conservation Service (NRCS) in effort to provide a consistent and objective method for relating rainfall to runoff, primarily for agricultural purposes in the mid-western United States (USDA NEH-4, 1985). The application of the CN method has since been extended beyond the original intent of use within small agricultural watersheds. In actual use by government agencies and practicing engineers, most CNs are drawn from agency tables derived from CNs originally developed for Midwestern agriculture or from consensus tables agreed upon

for local usage (Hawkins *et al.*, 2006). The Curve Number method is highly sensitive to the selected CN. If the CN is selected from tables derived for another region, it is possible that the CN and predicted runoff volume would be inaccurate given the local conditions, and would result in over or under-design. For the CN method to provide users with consistent data, local calibration of CN values is needed.

CN calibration is ideally performed by direct measurement using observed rain and stream gauge data. While this method is ideal, it is difficult to accomplish. Retrieving enough usable observed rain and stream gauge data in order to produce an accurate calibration is time consuming and often not possible.

Existing networks of National Oceanic and Atmospheric Administration (NOAA) rain gauges and United States Geologic Survey (USGS) flow stations can be used for CN calibration. However, in many cases, the project site is not located near any rain or stream gauges. Thus direct CN calibration measures are often unachievable over a wide area, such as the state of Utah. A more feasible method of calibration is necessary for broad use of the CN method in ungauged regions.

Due to the unavailability and difficulty in obtaining observed data, the USGS has developed regional regression equations for estimating flood magnitude and frequency at ungauged sites. These regression equations are used to interpolate or transfer flood characteristics from gauged to ungauged sites through the use of watershed and climatic characteristics including area, channel characteristics, elevation, and mean annual precipitation among others. These equations have been developed on a state-by-state basis for hydrologically similar regions. In 1994, the USGS developed a computer program called the National Flood Frequency Program (NFF). All of the regression

equations were compiled into a database for use in this program, which is often used by engineers and accepted by the Federal Highway Administration (FHWA) and Federal Emergency Management Agency (FEMA) for planning and design applications. Although the CNs calibrated using this method are influenced by the assumptions made in the development of the regression equations, use of the USGS regression equations is a viable and practical substitute for obtaining real data given its foundation in historical gauged data.

In order to perform CN calibration, the rainfall depth and associated runoff depth must be known for a given storm event in a given watershed. For the method of calibration developed in this study, the peak runoff flow predicted using the USGS regression equations are indexed by return-period and used in conjunction with a dimensionless hydrograph to compute the total runoff volume for each return-period. The runoff volume can be converted to an average runoff depth by dividing by the area of the watershed in question. Rainfall grids provided by NOAA are used to estimate the precipitation depth by return-period. Given the rainfall and runoff depth, a site-specific CN can then be calibrated for each return-period.

The Utah Department of Transportation (UDOT) is frequently involved with projects and hydrological studies throughout the state of Utah. UDOT uses the CN method for the design of structures in non-urban areas. A request from UDOT for a locally calibrated CN table was the impetus of this research. Calibration of the CN table currently used by UDOT has been performed using the regression equations and will be used as an illustration of this process.

The purpose of this study is to develop a method of CN calibration that can be easily applied to regions where gauged data are unavailable using the USGS regression equations. In this study, regression equation calibration was performed for nearly 60 watersheds throughout the state of Utah. The calibration results from these watersheds have been investigated for regional trends and have been compared to calibration results that were obtained using measured rain and stream gauge data. While much of this research has been focused on watersheds within the state of Utah, this method can be applied to any location where flood regionalization studies have been performed and regression equations are available. Calibration results obtained through the use of the USGS regression equations appear consistent with results obtained from gauged data. CN calibration performed through using the regression equations can potentially enable engineers and other organizations to more easily calibrate CN tables for local conditions and improve the adequacy of their designs.

## 2 Background

The CN method, USGS regression equations, design hydrographs and NOAA rainfall grids are all quite commonly used in engineering practice. CN calibration using the regression equations combines many of these engineering tools. To clarify these concepts, a background of the CN method, the USGS regression equations and the NFF dimensionless hydrograph is given.

#### 2.1 Curve Number Method

There are many factors that influence the infiltration and runoff volume within a watershed including the intensity of the storm, the volume of precipitation, the land use (i.e. forested, residential, etc.), soil type (i.e. clay, sand, etc.) and initial moisture conditions (i.e. wet, dry, etc.). The CN method is empirical in nature and based on land use, soil type and antecedent moisture conditions. The following is an overview and explanation of the Curve Number method.

#### 2.1.1 Curve Number Method Derivation

At the beginning of a rainfall event, the rainfall intensity is usually less than the rate at which water is stored, meaning all the precipitation is absorbed into the soil. However, as the storage is filled, and the soil and vegetation become saturated, there is

less capacity for storage or absorption. The precipitation is in excess and begins to "runoff" the land surface into streams, rivers and lakes. If the land is well forested, the rainfall is readily absorbed by the vegetation. However, if the land is paved as it is in a residential area, there is little absorption and nearly all precipitation becomes direct runoff.

Water flows more freely through sand than through clay thus precipitation is absorbed by sand but flows across the top of a clay surface. If the soil is initially saturated from a previous event, there is less storage capacity and most of the precipitation from the event following will become runoff (Wanielista *et al.*, 1997).

In general terms runoff is represented by the following equation:

$$R = P - S \tag{2-1}$$

where R = rainfall excess or runoff (in)

P= rainfall volume (in)

S= storage volume including initial abstraction and infiltration (in)

Initial abstraction is water intercepted by vegetation and stored in surface depressions (Wanielista *et al.*, 1997). At saturation, the rate of rainfall excess is equal to the intensity of precipitation. A proportional relationship can be developed as:

$$\frac{S}{S'} = \frac{R}{P} \tag{2-2}$$

The relationship of these factors is illustrated below in Figure 2-1.

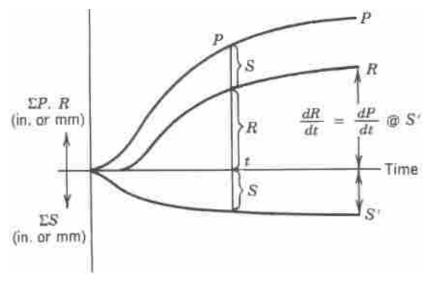


Figure 2-1: Time Variability of Hydrologic Events (Wanielista et al., 1997)

Additional work done by the SCS identified an empirical relationship between the initial abstraction ( $I_a$ ) and storage and developed an equation where  $I_a$  is defined as 0.2S'. However, abstraction values vary for different soil types. The empirical equations developed by the SCS to estimate maximum water storage capacity and rainfall excess are estimated using the equations below.

$$S' = \frac{1000}{CN} - 10 \quad \text{(in)}$$
 (2-3)

$$R = \frac{(P - 0.2S')^2}{(P + 0.8S')}$$
 (in) (2-4)

if 
$$P \ge 0.2S'$$
 otherwise  $R = 0$ 

A graph of this equation is shown in Figure 2-2 below:

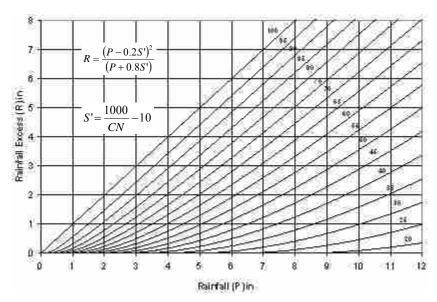


Figure 2-2: Curve Number on Rainfall vs. Rainfall Excess Plot (Wanielista et al., 1997)

## 2.1.2 Curve Number Affecting Factors

Three main factors that affect the CN value are antecedent moisture condition, soil type, land use.

### 2.1.2.1 Antecedent Moisture Conditions

The NRCS has established three antecedent moisture conditions. These conditions are as follows:

## • Condition 1:

A condition of drainage basin soils where the soils are dry but not to wilting point

## • Condition 2:

The average case

## • Condition 3:

When heavy rainfall or light rainfall with low temperatures have occurred, producing high runoff potential.

Most CN tables are developed for Condition 2, the average case. However, there are adjustment factors available for Conditions 1 and 3 (Wanielista *et al.* 1997).

## **2.1.2.2** Soil Type

The soil type is divided into four classes, A through D based on the ability of water to pass through the soil. Table 2-1 identifies the rate (inches per hour) at which water passes through each soil group.

Table 2-1: Soil Type Classifications Based on Transmission Rate of Water (Hawkins et al., 2006)

Group	Soil Type	Tranmission rate (in/hr)
Α	Sand, loamy sand, sandy loam	greater than 0.30
В	Silt loam or loam	0.15 to 0.30
С	Sandy clay loam	0.05 to 0.15
D	Clay	0 to 0.05

#### 2.1.2.3 Land Use

As stated previously, land that is well forested readily absorbs rainfall whereas in areas that are covered with impervious surfaces such as pavement, nearly all precipitation becomes direct runoff. The variation in the CN due to land use can be seen in the CN table shown below in Table 2-2 which was developed for antecedent moisture condition II. Variation in the CN due to soil type is found under the "Hydrologic Soil Group" column. The hydrologic condition is based on a combination of factors that affect infiltration and runoff including the density of vegetated areas, amount of year-round vegetative cover, degree of surface hardness etc. "Good" hydrologic conditions encourage average and better than average infiltration and tend to decrease runoff. "Poor" hydrologic conditions tend to increase runoff (USDA NEH-4, 1985).

Table 2-2: Runoff Curve Numbers (USDA NEH-4, 1985)

Cover description	***************************************	· · · · · Hydrologic soil group · · · · ·			
cover(ype	hydrologic condition <sup>2/</sup>	A <sup>3/</sup>	В	C	D
Herbaceous—mixture of grass, weeds and low-growing	Poor		80	87	93
brush, with brush the minor element	Fair		71	81	.89
	Good		62	74	85
Oak-aspen—mountain brush mixture of oak brush, aspen,	Poor		66	74	79
mountain mahogany, bitter brush, maple, and other brush	Fair		48	57	63
	Good		30	41	48
Pinyon-juniper—pinyon, juniper, or both; grass understory	Poor		75	85	89
	Fair		58	73	80
	Good		41	61	71
Sage-grass—sage with an understory of grass	Poor		67	80	85
	Fair		51	63	70
	Good		35	47	55
Desert shrub—major plants include saltbush, greasewood,	Poor	63	77	85	.88
creosotebush, blackbrush, bursage, paloverde, mesquite,	Fair	55	72	81	86
and cactus	Good	49	68	79	84

V Average runoff condition, and  $I_n = 0.2s$ . For range in humid regions, use table 9-1.

<sup>27</sup> Poor; <30% ground cover (litter, grass, and brush overstory).</p>
Fair: 30 to 70% ground cover.

Fair: 30 to 70% ground cover. Good: >70% ground cover.

<sup>3</sup>º Curve numbers for group A have been developed only for desert shrub.

CN values are very dependent upon local conditions. "Forested" areas in Tennessee can be quite different than what might be considered "forested" in Utah. Thus, local conditions should be taken into careful consideration when using the CN method in order to avoid inaccurate runoff estimations.

### 2.1.3 CN Back-Calculation Derivation

Typically, the CN method is used to estimate runoff, however, if the rainfall and runoff is known, a CN for a particular watershed and storm event can be back-calculated. The derivation of this procedure is as follows. Equation 2-3 is rearranged in order to solve for CN:

$$S' = \frac{1000}{CN} - 10\tag{2-3}$$

where 
$$CN = \frac{1000}{S'+10}$$

Equation 2-4 is also rearranged and then inserted into the quadratic equation to solve for S' in terms of R and P as follows:

$$R = \frac{(P - 0.2S')^2}{(P + 0.8S')}$$
 (2-4)

Multiplying each side by (P + 0.8S') results in:

$$R(P+0.8S') = (P-0.2S')^2$$
.

All terms are then distributed as seen below.

$$PR + 0.8S'R = P^2 - 0.4S'P + 0.04S'^2$$

The equation is set to zero be moving all terms to the right side.

$$0 = 0.04S'^2 - 0.4S'P - 0.8S'R + P^2 - PR$$

The equation is then multiplied by 25 for ease of use, as seen below.

$$0 = S'^2 - 10S'P - 20S'R + 25P^2 - 25PR$$

Factoring out the 10S' simplifies the equation for insertion into the quadratic equation.

$$0 = S'^2 - 10S'(P + 2R) + 25P^2 - 25PR$$

The general form of the quadratic equation is:

$$y = \frac{-b + or - \sqrt{b^2 - 4ac}}{2a}$$
 (2-5)

where 
$$0 = ay^2 + by + c.$$

In this case,

$$y = S'$$

$$a = 1$$

$$b = -10P - 20R$$

$$c = 25P^{2} - 25PR$$

which results in

$$S' = \frac{-(-10P - 20R) \pm \sqrt{(-10P - 20R)^2 - 4(1)(25P^2 - 25PR)}}{2(1)}.$$

This equation can then be reduced to

$$S' = \frac{10P + 20R \pm \sqrt{100P^2 + 400PR + 400R^2 - 100P^2 + 100PR}}{2},$$

and ultimately to

$$S' = 5(P+2R) \pm 5\sqrt{5PR+4R^2} \ . \tag{2-6}$$

After having solved for S' in terms of R and P, we can solve for CN in terms of R and P: From Equation 2-3,

$$CN = \frac{1000}{S'+10}$$

Substituting in Equation 2-6 for S' yields:

$$CN = \frac{1000}{5(P+2R) - 5\sqrt{5PR + 4R^2} + 10}$$

which is simplified to:

$$CN = \frac{200}{P + 2R - \sqrt{5PR + 4R^2} + 2} \tag{2-7}$$

The negative root is selected to obtain the correct solution for Equation 2-7 (Curtis et. al. 2, 1983). This equation can be used to back-calculate a CN when the rainfall and the resulting runoff for a given storm are known. The CN calibration results of this study were obtained using this equation.

## 2.2 USGS Regional Regression Equations

The USGS has developed regional regression equations which are used to transfer flood characteristics from gauged to ungauged sites. These equations aid engineers in the design of projects in where gauged rainfall and runoff data is unavailable. The regression equations have been developed on a regional basis and were compiled into a database which is used in a computer program called the National Flood Frequency Program (NFF) (Reis et al., 2002). The concept of expanding the utility of gauged site data for use at ungauged locations with similar physiographic and climatic characteristics is not new, and several methods have been examined and tested during the past 50 years. The method of choice for the past 30 years has been statistically based regional equations that predict peak stream flow. This method involves a division of the study area into regions of similar physiographic and climatic characteristics.

#### 2.2.1 Flood Regionalization

The testing of regional homogeneity is a critical part of flood regionalization procedures due to the substantial variability of flood characteristics that may exist between regions with differences in climate, topography, and geology. Multiple-linear regression techniques are applied to determine coefficients for statistically significant predictors of peak stream flow i.e. basin area, elevation, and mean annual precipitation.

Upon determining the level of influence these variables have on peak stream flow, regression equations can then be derived accordingly (USGS 1999).

#### 2.2.2 Return-Period

For every region, regression models or equations are developed for each return period or recurrence interval flow such as the 2-, 5-, 10-, 25-, 50-, 100-, 200-, and 500-year peak flow. A return-period or recurrence-interval is an estimate of the likelihood of events. It is a statistical measurement indicating the average recurrence interval over an extended period of time. A specified return-period is usually required for risk analysis in order to design structures so that they are capable of withstanding an event of a certain return-period and its associated intensity (USGS 1999).

### 2.2.3 Use and Development of the NFF Dimensionless Hydrograph

A runoff hydrograph is a graph of surface water discharge over time. A runoff hydrograph has three main parts: a rising limb, peak discharge and a falling limb. The area under the hydrograph represents the total volume of runoff produced by a given storm event. The variation of hydrograph shape is due to varying storm intensity over time and unique watershed characteristics (Wanielista *et al.* 1997). A hydrograph for a single storm event is shown below in Figure 2-3.

The hydrograph in Figure 2-3 features a double peak. Multiple peaks are common in hydrographs that depict a naturally occurring storm event. Often the intensity of the storm and, consequently, the runoff will vary over the duration of the storm.

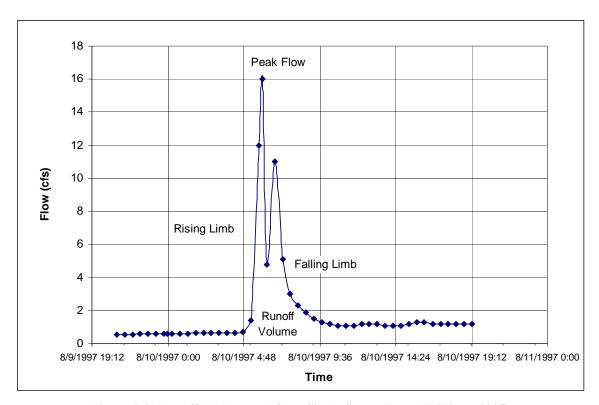


Figure 2-3: Runoff Hydrograph for a Single Storm Event (Williams 2005)

## 2.2.3.1 NFF Dimensionless Hydrograph Use

Due to the difficulty in obtaining observed rainfall and runoff data for a watershed, the use of "synthetic" or design hydrographs has become a common engineering practice. Design hydrographs are used to simulate flood hydrographs in areas where gauged data are unavailable. A design hydrograph does not represent runoff from an actual storm, but is developed using a dimensionless hydrograph. A dimensionless hydrograph is an average of several hydrographs obtained from gauged data in a particular region. The dimensionless hydrograph is referred to as "dimensionless" because the ordinates are scalars with no units associated with them. The development of these ordinates is explained below. The NFF dimensionless hydrograph is shown in Figure 2-4 below

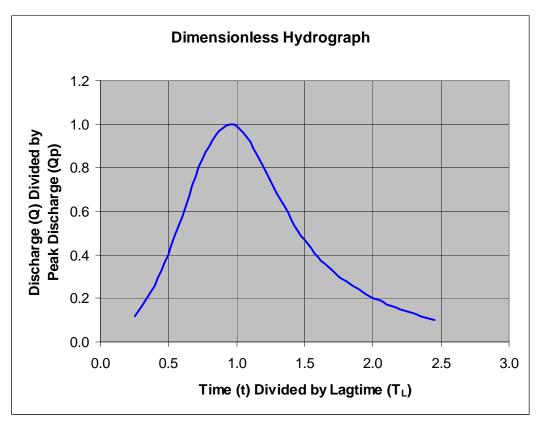


Figure 2-4: NFF Dimensionless Hydrograph

Three elements are needed in order to develop an NFF design hydrograph including peak discharge, basin lag time, and the dimensionless hydrograph ordinates. The peak flow is calculated in NFF for a specified return-period using the USGS regression equations. Basin lag time is calculated from empirical methods based on the general characteristics of runoff lag time including watershed hydraulic length, average watershed slope and storage potential within the watershed. Theoretically, lag time is defined as the time between the center of mass of rainfall excess to the time of peak runoff as seen in Figure 2-5. Although there are several methods of calculating lag time, most produce similar results. In previous research, several methods for estimating lag time were tested, and it was determined that the Denver method was suitable for use in this study (Dyer 2006).

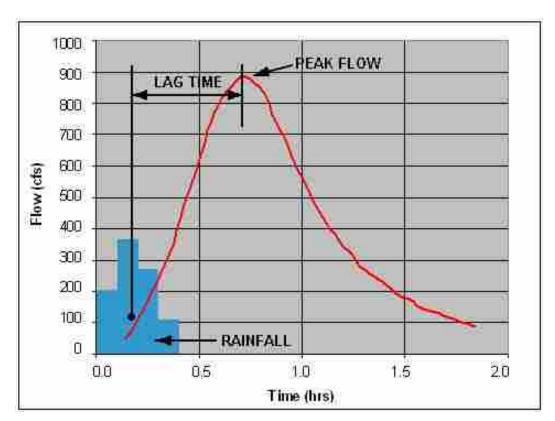


Figure 2-5: Graph of Definition of Lag Time

The ordinates of the NFF dimensionless hydrograph are used in combination with the calculated peak flow and lag time to compute the design hydrograph (FEMA 2007). The "x" ordinates ( $t/T_L$ ) of the dimensionless hydrograph are multiplied by the lag time in order to calculate the "x" ordinates of the design hydrograph. Likewise, the "y" ordinates (Q/Qp) of the dimensionless hydrograph are multiplied by the peak flow to calculate the "y" ordinates of the design hydrograph. Two design hydrographs developed for two different watersheds using the NFF dimensionless hydrograph are shown below in Figure 2-6. As seen in Figure 2-6, as the lag time increases the hydrograph is extended over a longer period of time.

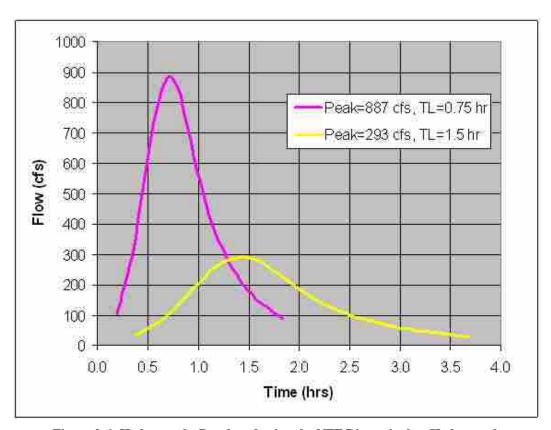


Figure 2-6: Hydrographs Developed using the NFF Dimensionless Hydrograph

Figure 2-7 depicts design hydrographs of multiple return-periods for a given watershed. The lag time is constant at each return-period since each hydrograph was calculated for the same watershed. However, as anticipated, the peak of the hydrograph increases with increasing return-period. Unlike the hydrograph shown in Figure 2-3, design hydrographs produced by NFF feature only a single peak. Although design hydrographs do not depict actual storm events, the design hydrographs produced by NFF can be used to estimate the volume of flow for each return-period, which is useful for some design purposes (Reis *et al.*, 2002). The use of NFF design hydrographs is particularly appropriate for CN calibration since the volume of runoff or area underneath the hydrograph is of key importance, and timing and shape of the runoff hydrograph is not significant.

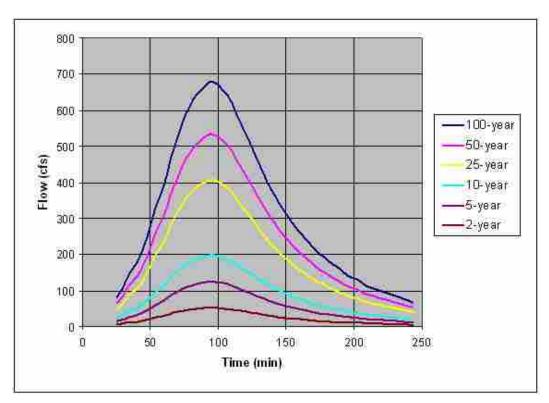


Figure 2-7: Hydrograph by Return-Period Developed using NFF

## 2.2.3.2 NFF Dimensionless Hydrograph Development

The NFF dimensionless hydrograph is based on a study performed by the USGS in cooperation with the Georgia Department of Transportation (Reis *et al.*, 2002). The dimensionless hydrograph was developed from more than 350 observed flood hydrographs and over a hundred gauging stations in both rural and urban areas of Georgia. Only simple (i.e. single peak) flood hydrographs that resulted from relatively uniform rainfall were selected for the study (Inman 1987).

Unit hydrographs as well as lag time were computed for several storms at each station. Unit hydrographs with inconsistent shapes were eliminated from the study. An average of the unit hydrographs was computed for each station by aligning the peaks and

averaging each ordinate of discharge. The timing of the average hydrograph was determined by averaging the lag times calculated for the individual storms (Inman 1987).

Several different durations (1/4. 1/3, 1/2 and 3/4 of the lag time) were applied to the average unit hydrographs. These hydrographs were then transformed into dimensionless hydrographs by dividing the time at each ordinate by the lag time and discharge at each ordinate by the peak discharge. The dimensionless hydrographs were grouped into regions of similar physiographic and climatic characteristics (i.e. rural regions, urban regions, regions with similar vegetation, etc). The dimensionless hydrographs were then averaged with other dimensionless hydrographs developed for the same duration at stations within the same region. Figure 2-8 is a plot of the average 1/2 lag time duration dimensionless hydrograph for Georgia's Region 1. The range of the data from the 16 stations from which it was computed is also plotted (Inman 1987).

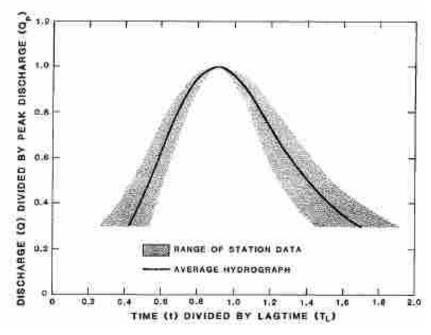


Figure 2-8: Region 1 Dimensionless Hydrograph for the ½ Lag Time Duration (Inman 1987)

Based on the regional average hydrographs, a single hydrograph was developed for application within both rural and urban areas over the entire state. Several statistical tests indicated that the ½ lag time duration was the best fit for the majority of the statewide data. To validate the dimensionless hydrograph, the dimensionless hydrograph was applied to 138 other observed hydrographs not used in its development (Inman 1987). A plot of the Georgia statewide dimensionless hydrograph method compared to other dimensionless hydrographs, including the SCS and Stricker-Sauer methods can be seen in Figure 2-9 below.

The dimensionless hydrograph developed for the state of Georgia has been adopted for use in the NFF program to compute design hydrographs for the entire nation. Except in some relatively flat topography, and slow runoff areas, the dimensionless hydrograph is applied nationwide with reasonable accuracy, particularly when it is used to estimate average runoff volume (Reis *et al.* 2002).

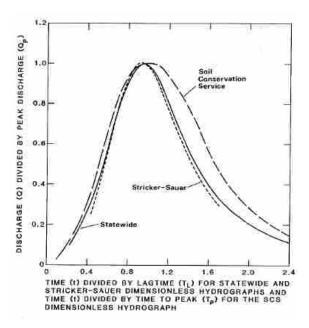


Figure 2-9: Georgia Statewide (NFF) Dimensionless Hydrograph Compared to the SCS and Stricker-Sauer Methods

## 2.2.4 Utah USGS Regression Equations

The state of Utah is located within a regional flood study area that encompasses the arid lands of the southwestern United States. The study area is divided into 16 hydrologic flood regions, of which 7 include portions of Utah. A map of these regions is shown in Figure 2-11 (USGS 1999). Within Utah, regions with an elevation greater than a specified threshold are considered to be in Region 1 (see Figure 2-10). While the use of regression equations is a convenient method for runoff estimation, there are limitations in their applicability. Table 2-3 summarizes these limitations by region. The equations developed for the state of Utah were developed specifically for non-urban regions. USGS also developed equations for urban areas that can be applied nationwide. For this study, however, only the Utah USGS regression equations were used, and are thus, only applicable in the rural areas of Utah.

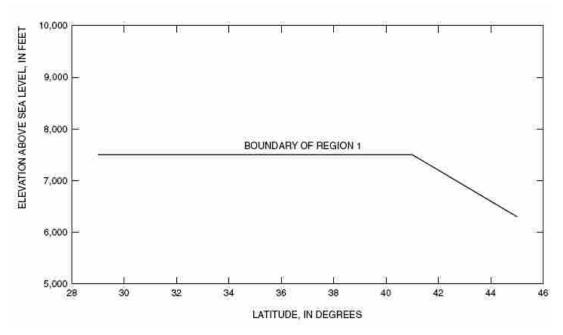


Figure 2-10: Utah NFF Region 1 Elevation Threshold (USGS 1999)

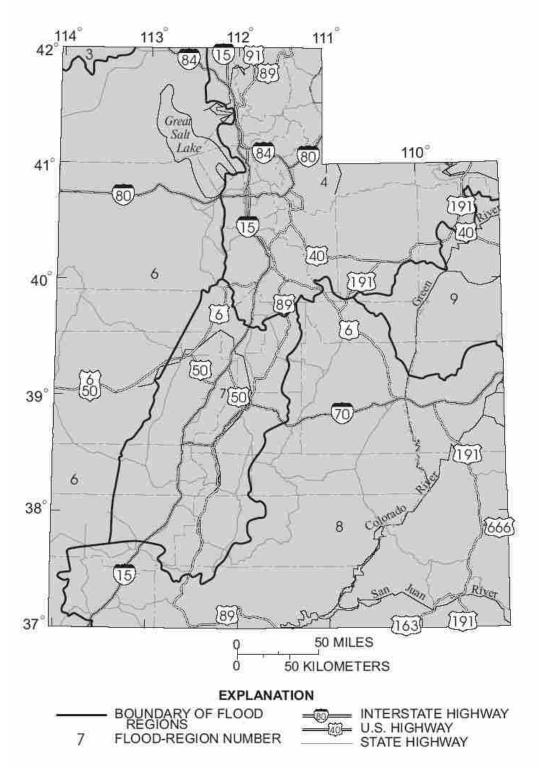


Figure 2-11: NFF Regions in Utah (USGS 1999)

Table 2-3: Limitations in Regression Equation Applicability (USGS 1999)

Hydrologic study region	Drainage area, in square miles <sup>1</sup>	Mean basin elevation, in feet above sea level <sup>2</sup>	Mean annual precipitation, in inches
Region I	$0.30 - 1.060^{1}$	ī.	10-45
Region 3	2.35-1,450 <sup>1</sup>	##:	10-46
Region 4	$2.08 - 1,230^{1}$	5,740-10,700	<del>50</del>
Region 6	.20-1,670 <sup>1</sup>	<del>98</del> 1	966
Region 7	5.58-3401	6,070-9,560	900
Region 8	.20-1,990 <sup>1</sup>	4,300-10,200	22
Region 9	4.00-1.180 <sup>1</sup>	6,550-10,060	

<sup>&</sup>lt;sup>1</sup>For best results, applications should be limited to basins of less than 200 square miles.

The first multiple-linear regression study of regional flood frequency for Utah was completed in 1971. Since 1971, more than five multiple-linear regression studies have been completed (Kenney *et al.* 2007). The regression equations used in the calibration of the Utah Department of Transportation CN table performed in this study are the most recent equations as of 1997. These equations can be seen in Appendix D. Studies for the next generation of improved regression equations for the state of Utah were recently completed by the USGS in October of 2007. These equations were unavailable during the development of this research. However, the methods set forth in this research for CN calibration are still applicable for the latest version of the regression equations. This method of CN calibration would be applicable within any state where flood regionalization studies have been completed by using the equations appropriate for the region of interest.

<sup>2</sup>NGVD of 1929.

# 3 Preceding Research

The concept of CN calibration is not new, and has been researched by many individuals and organizations. CN calibration performed using USGS regression equations is an extension of previous CN calibration research done by Brigham Young University (BYU) faculty and students (Williams 2005, Dyer 2006). The following is a summary of the background research and concepts which have been the foundation and support of the development of CN calibration using USGS regression equations.

## 3.1 CN Calibration using Measured Data

CN calibration accomplished using measured data is preferred over other methods of calibration since the results are more accurate and site-specific. In 1983, students from Utah State University (USU) under the direction of Dr. Richard H. Hawkins compiled and published "A Catalog of Intermountain Watershed Curve Numbers" (Curtis *et. al* 1983). The data used for the watersheds located in Utah were taken from a Bureau of Land Management (BLM) study. Calibration was done using two methods. The first was the Rallison-Cronshey method which is the same as previously outlined in 2.1.3. The second was an optimized least squares method for which the determination of S' is done iteratively. A more detailed explanation of this methodology is available in the catalog

compiled by USU (Curtis *et. al* 1983). CN calibration results from this catalog for watersheds in the Price, Utah area are shown below in Table 3-1 through Table 3-3.

Table 3-1: CN Values Calibrated by USU for Coal Creek Tributary (Curtis et al., 1983)

Location	lat 39.33.40	55	Area: 0.41 mr/2						
Storm	Date	Rainfall		Runoff		Event CN			
1	10/13/1981	0.30	in.	0.039	in	94.2			
2	10/3/1981	0.20	in	0.004	in	93.3			
3	10/16/1981	0.75	in	0.075	in	85.1			
1	10/4/1981	0.40	in	0.001	in	84.8			
5	10/12/1981	0,65	in	0,039	in	84.6			
	Minimum	0.20	in	0.001	in	84.6			
	Average	0.46	in	0.032	in	88.4			
	Maximum	0,75		0,075	in	94.2			
	Smd. Dev.	0.23		0.030		4.9			

Table 3-2: CN Values Calibrated by USU for Solider Creek Tributary (Curtis et al., 1983)

	Cr <del>ook</del> Tributa									
Location: lat 39.33.44 long, 110,39.20 Area: 1.25 r										
Storm	Date	Rainfall		Runoff		Event Ch				
1	10/13/1981	0.15	in	0.037	in	98.0				
2	10/11/1981	0.26	in	0.067	īn:	96.6				
3	10/15/1981	0.40	in	0.076	in	93.7				
4.	10/3/1981	0.40	in	0.072	in	93.5				
5	10/4/1981	0.20	in	0.001	in	92.2				
6	10/16/1981	0,66	in	0.120	in	90.1				
	Minimum	0.15	in	0.001	in	90.1				
	Average	0.34	in	0.062	in	94.0				
	Maximum	0.65		0.120	īn:	98.0				
	Stnd. Dev.	0.18		0.040		2.9				

Table 3-3: CN Values Calibrated by USU for Wattis Branch (Curtis et al., 1983)

Wattis B	ranch										
Location: lat. 39.31.30, long. 110.52.1 Area: 4.9 mi <sup>2</sup>											
Storm	Date	Rainfall		Runoff		<b>Event CN</b>					
Ť	10/11/1981	0.10	in:	0.009	in	97.6					
2	10/17/1981	0.17	in	0.026	in	96.9					
3	9/5/1981	0.17	in	0.010	in	95.4					
4	10/3/1981	0.20	in	0.010	in	94.4					
5	10/11/1981	0.50	in	0,032	in	88,0					
6	10/16/1981	0,60	in	0.044	in	86.4					
	Minimum	0.10	ın.	0.009	in	86.4					
	Average	0.29	in	0.022	in	93.1					
	Maximum	0.60		0.044	in	97.6					
	Stnd. Dev.	0.21		0.015		4.8					

Results for these watersheds and storm events resulted in relatively high Curve Numbers. Given that these storm events produced rainfall less than that of a 2-year return-period event, high CN values would be expected since CN values generally decrease with increasing return-period (Hawkins 2007). The results presented in this section will be compared to results for the same watersheds obtained using the USGS regression equations in the following chapter.

Although using measured data produces more accurate and site-specific results than other calibration methods, this method of CN calibration has some limitations. Setting up enough gauges to be able to accurately interpret the rainfall-runoff relationship can be a challenge. The data collected in the BLM study, as seen in Table 3-1 through Table 3-3, illustrates that encountering storm events of a significant size is not an easy task. Measurement must take place over an extended period of time in order to be able to come across a storm large enough to clearly observe the runoff response to a storm event and to provide reliable results. Predicting the frequency and magnitude of flooding would also not be possible if measured data were only collected over a short period of time. Use

of measured data can generate more precise CN calibrations but the collection of quality data is difficult and uneconomical if CN calibrations are needed over large areas.

#### 3.2 CN Calibration using Historical Data

Although CN calibration using measured data poses many challenges, there are, however good alternatives to direct measurement. The USGS has numerous stream gauges all over the country as well as within the state of Utah. Decades of historical data are available for each stream gauge. The National Climatic Data Center (NCDC) maintained by NOAA likewise has extensive historical records for numerous rain gauges. Many of these historical records are available via the internet while others are available upon request. As discussed in 2.1.3, if the precipitation (P) and the runoff (R) are known, a CN can then be back-calculated and calibrated for watersheds where the rain and stream gauge data are available. If enough historical data were available, CN calibration on a large scale would be possible using these historical records.

Joel Williams (Williams 2005) of Brigham Young University compiled a database of precipitation and stream flow data for various watersheds in the state of Utah in order to determine the feasibility of developing local CN calibrations with available historical USGS and NOAA data. Several factors must be taken into consideration when using historical data such as:

- Close proximity of the stream and rain gauges which is necessary in order to observe a more direct rainfall-runoff relationship
- Only streams with very few diversions can be used for calibration so as to avoid alteration in flow readings

- Only records occurring during the summer and fall months (July-September)
   can be used so as to avoid the effects of snowmelt runoff
- Selected storm events must be of great enough significance to produce a noticeable rise in the stream volume (i.e. 0.5 inches or greater)
- Only precipitation records that occurred while the stream gauges were in operation can be used and vice versa so a direct rainfall-runoff relationship can be observed.

Williams was able to locate overlapping precipitation and stream gauge records that met this criterion for a total of 40 rain-stream gauge pairs. The actual hydrograph and precipitation data taken from the historical records for one of the gauge pairs is shown in Figure 3-1 below.

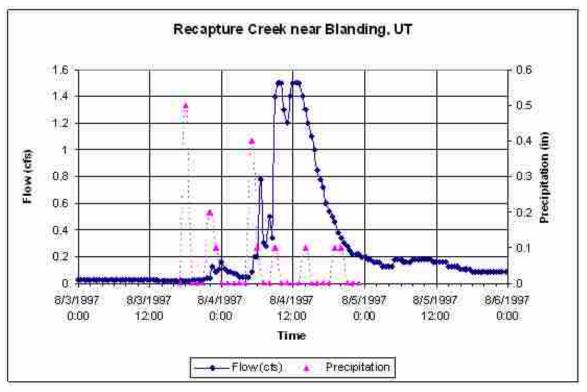


Figure 3-1: Stream and Precipitation Graphs (Williams 2005)

The CN calibration process used by Williams is as follows:

- The watershed was delineated using Watershed Modeling System (WMS)
   (Nelson 2006), a software program developed by Brigham Young University
- Collected precipitation data was imported into WMS
- A hydrologic runoff model using the SCS unit hydrograph was created from the delineated watershed and precipitation data
- The hydrologic runoff model was then used to determine runoff volume which was calculated from the computed hydrograph
- The input CN was iterated until the predicted runoff volume from the WMS model matched the actual runoff volume from the hydrograph

The results of the CN calibration for these gauge pairs varied with each watershed. When compared to the composite CNs calculated using the CN table currently used by UDOT (see Appendix A), in some cases, the calibrated CN increased while in other cases the CN decreased. While much of the data appeared reasonable there were instances where the calibration produced CN values exceeding 100 or where half the amount of rainfall resulted in twice the runoff from one storm event to another within the same watershed. These discrepancies could be due to gauge malfunction or the fact that the rain and stream gauge were perhaps not close enough to ensure that the same storm event was being measured by both gauges. If a watershed were gauged specifically for the purpose of CN calibration, data overlap, gauge malfunction and proximity issues would not be a concern.

While Williams' research indicates that it is possible to back-calculate a CN from historical data, this method obviously cannot be used in regions where no gauges are

present. Historical data gauged by others can be unreliable and result in inaccurate calibrations. Although rain and stream gauge data is easily accessible on the internet, painstaking effort was required to first, find concurrent rain and stream gauge data; second, locate gauges in close enough proximity to ensure measurement of the same event; and lastly, find watersheds with available data that meet all other criterion. Use of data measured by others appears convenient but reliable CN calibrations are difficult to obtain using this method.

## 3.3 Data Catalog and Script for CN Calibration

Hydrological study methods have changed dramatically over the last few decades. Much of the processes of today are done digitally. Due to the digitization of hydrological studies, there is an abundance of digital data that are available through government and other agencies via the internet. Digital data that is useful for CN calibration includes:

- Digital Elevation Model (DEM) (<a href="http://seamless.usgs.gov/">http://seamless.usgs.gov/</a>)
- Land Use Coverage (http://www.webgis.com/lulcdata.html)
- Soil Type Coverage
   (<u>http://www.epa.gov/waterscience/ftp/basins/gis\_data/huc/\_</u>)
- Precipitation Frequency Data (<a href="http://hdsc.nws.noaa.gov/hdsc/pfds/">http://hdsc.nws.noaa.gov/hdsc/pfds/</a>)
- Topographic Maps (http://terraserver.homeadvisor.msn.com/)

Though all of these data are readily available and accessible via the internet, the retrieving and importation of data into software programs can be a time intensive process. If the data are to be used on a regular basis, the compilation of the data into a catalog and

the use of a script or simple computer code to locate and process the data can be helpful in streamlining design processes.

Shane Dyer (Dyer 2006) of BYU compiled DEMs, land use and soil type coverages, rainfall grids for precipitation estimation and topographic maps for the whole state of Utah into a computer catalog or database. The catalog compiled by USU, mentioned previously, was a collection of calibrated CNs for user reference. The catalog developed by Dyer was a collection of digital data that could be used to calibrate CNs, and was specifically made for the purpose of calibrating the UDOT CN table using WMS. The NFF program is integrated within WMS. All of the digitized data can be overlaid in WMS and processed for CN calibration using the USGS regression equations. The step-by-step process is available in Appendix B.

Dyer (Dyer 2006) also created a script for the automation of the calibration procedure. The script opens WMS and inputs map files containing the drainage basin area, land use and soil type coverages. With this data, the NFF database which contains the USGS regression equations can then be used within WMS to calculate peak flow and to develop a design hydrograph for each return-period. Results for each watershed are exported to a text file (Dyer 2006). The script output text file can then be imported into an Excel spreadsheet for further calculations. An example spreadsheet is shown in Figure 3-2. The data contained in the spreadsheet is as follows:

- **FILENAME**: Filename of the WMS map file used in the script.
- **UDOT**: Composite CN for the watershed calculated using the CN table currently used by UDOT for comparison purposes.

- **CLASS**: Composite CN for the watershed calculated using the CN table researched by BYU students for comparison purposes.
- **P2 through P100**: Precipitation for each return-period from NOAA Atlas 14 rainfall grids. The precipitation as well as the volume, runoff and calibrated CNs, are indexed by return-period as indicated by the numbers 2 through 100.
- **V2 through V100**: Runoff volume for each return-period calculated for each return-period using the NFF design hydrograph. Again, numbers 2 through 100 indicate the 2 through 100-year return-period.
- Area: Watershed area calculated in WMS.
- **R2 through R100**: Average runoff depth calculated by dividing the runoff volume (V2 through V100) by the watershed area.
- **CN 2 through CN 100**: Back-calculated CN for each return-period using P and R values for the respective return-period in Equation 2-7:

$$CN = \frac{200}{P + 2R - \sqrt{5PR + 4R^2} + 2} \tag{2-7}$$

The compilation of this database and development of the script would allow CN calibration using the USGS regression equations for particular watersheds to be performed quickly. However, much research as to the validity of this method was yet to be explored.

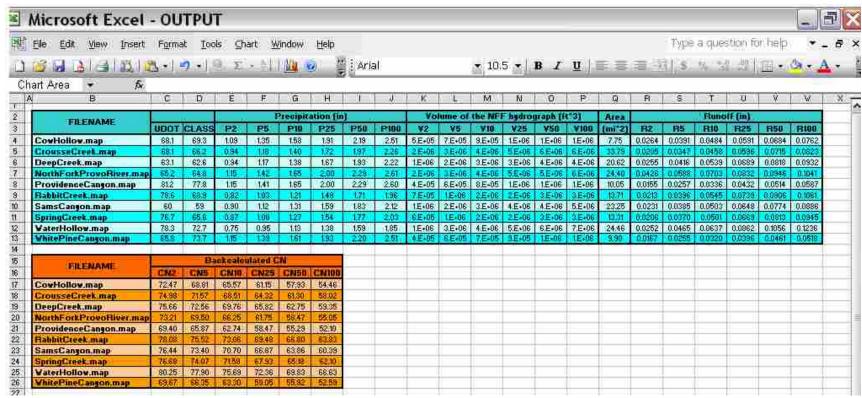


Figure 3-2: Script Output in Excel Spreadsheet

## 4 Methods and Procedures

Reliable measured data are difficult to obtain. Consistent CN calibrations are also consequently difficult to obtain. The USGS regression equations were developed for this reason: to provide good estimations of peak flow and runoff where gauged data is unavailable. Use of the USGS regression equations is pervasive throughout the engineering industry. The regression equations are also rooted in historical gauged data. All of these factors support the potential use of the regression equations in CN calibration.

UDOT currently uses the CN method in many of their design procedures. Along with other organizations within the industry, UDOT is unlikely to develop new methods of runoff estimation due to the familiarity and the widespread use of the CN method. UDOT does however desire to improve the accuracy and consistency of the method's use within the state of Utah by calibrating their current CN table for local conditions. Given that the direct measurement and collection of data for CN calibration can be such a demanding process, a more feasible method of calibration is necessary for the extensive (i.e. the whole state of Utah) CN calibrations needed by UDOT. UDOT also frequently uses USGS flood estimation for many of their projects. Considering the scale at which calibration is needed, and UDOT's previous experience with and use of the USGS regression equations in design, the use of USGS regression equations for CN calibration

would be an appropriate and practical method of obtaining data for use in CN calibration.

Thus, the UDOT CN table will be used as an illustration of this process.

In order to validate the use of the regression equations in CN calibration, regression equation calibration was applied to regions where gauged data has been obtained previously in the studies by Utah State University students and Joel Williams of BYU as discussed in 3.1 and 3.2 respectively.

## 4.1 Measured Data CN Calibration Comparison

The watersheds located in Utah that were researched by Utah State University (Curtis *et. al* 1983) were used for comparison in order to validate the use of regression equations in CN calibration. These three watersheds are all located near Price, Utah. In order to compare these methods, a DEM, land use shapefile and soil type shapefile for the region surrounding the basin outlet coordinates given in Table 3-1 through Table 3-3 were imported into WMS using the catalog compiled by Dyer (2006). The watersheds were then delineated and the resulting map files were saved and process using the script also developed by Dyer (Dyer 2006). The complete process utilized in WMS is outlined in Appendix B. With the availability of the catalog and script, only the first few steps of Appendix B are done manually. The remaining steps have been automated. The script output is then imported into WMS so that the CN values for the respective return-periods could be calculated. The output spreadsheet for the three watersheds (Coal Creek Tributary, Solider Creek Tributary and Wattis Branch) previously studied by Utah State University is shown below in Table 4-1.

Table 4-1: Script Output and Calibrated CNs for Utah State University Watersheds

FILENAME Presidentian II							0		Volume of the NFF hydrograph (#*3)							Runsit (in )					
111/2017/011/20	UDOT	CLASS	P2.	P.5	P10	P25	P50	P100	- V2	, vn	VIII)	V25	V30	V100	Imi _1	R2	R5	RID	R25	-R50	R100
Coal Creek	76.8	76.8	0.74	0:92	1.08	1_30	1.50	1.75	2:E+03	1.E+84	2:E+94	1.E+96	2.E+05	3.E+05	0.40	0.0027	0.0119	0.0263	0.1193	0.1873	0.2817
Solider Creek		77.	0.77	0.95	1.12	35	1.55	1,81	1.E+04	5.E+04	1.E+05	4.E+05	0.E+05	9.E+05	1.24	0.0048	0.0184	0.0372	0.1400	0.2087	0.2980
Wattis Branch	78.6	78.6	0.82	1.01	1,18	1.42	1.64	1,91	9 E+04	3.E+05	5.E+05	1.E+06	2.E+06	2.E+06	504	0.0079	0.0229	0.0404	0.1081	0.1495	0,2002
FILENAME	ENERGIGINAL ENGLISHMENT ON					1															
THE COUNTY IS	€II2:	CN5	CMIO	CH25	:CNS8	CN100:															
Coal Creek	75.56	73.76	72.75	76/12	76.07	75.54															
Solider Creek	76.79	74.28	73.28	76.41	76.01	75.14	i.														
Wattis Branch	75.44	73.67	72.29	73.07	71.65	69.74	1														

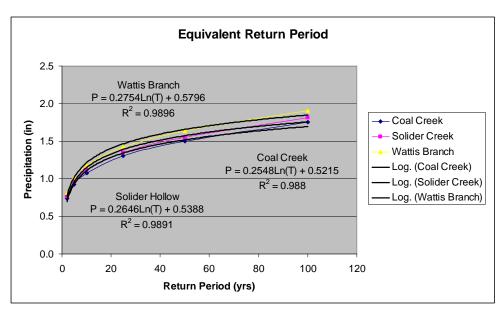


Figure 4-1: Equivalent Return-Period Equation Derivation for USU Watersheds

In order to compare the resulting CN calibrations, the equivalent return-period of the precipitation data used by USU was determined.. This was accomplished by first graphing the return-period against the precipitation values from the output file (P2 through P100) as seen in Figure 4-1. A trend line and an associated equation were developed for each watershed. The equations relate return-period (T) to precipitation depth (P). The equations can also be seen in Figure 4-1. The precipitation data from Table 3-1 through Table 3-3 were inserted into the respective equations to solve for a return-period (T). For example, the trend line equation for Coal Creek Tributary is:

$$P = 0.2548Ln(T) + 0.5215 (4-1)$$

Storm 1 for Coal Creek Tributary in Table 3-1, the rainfall was 0.3 inches. Inserting 0.3 into the above equation results in a return-period (T) of 0.43 years. This process was preformed for each storm evaluated by USU. The CN values from the output spreadsheet were then plotted against return-period. The CNs calibrated by USU for each storm were plotted on the same graph with the associated equivalent return-period previously calculated. The average of all the storms was also plotted with a line connecting to the data produced using the script and USGS regression equations to better illustrate their relationship. The results are shown Table 4-2 through Table 4-4 and Figure 4-2 through Figure 4-4 below.

**Table 4-2: Coal Creek Data Comparison** 

USU Co	USU Coal Creek Tributary Data												
Storm	CN	P (in)	T (yrs)										
1	94.17	0.3	0.42										
2	93.27	0.2	0.28										
3	85.15	0.75	2.45										
4	84.85	0.4	0.62										
5	84.62	0.65	1.66										
Avg	88.41	0.46	0.79										
BYU 2-year	75.56	0.741	2										

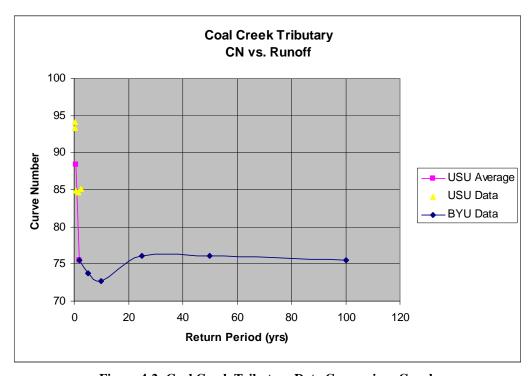


Figure 4-2: Coal Creek Tributary Data Comparison Graph

**Table 4-3: Solider Creek Comparison Data** 

USU Solider Creek Tributary Data												
Storm	CN	P (in)	T (yrs)									
1	97.96	0.15	0.23									
2	96.62	0.26	0.35									
3	93.74	0.40	0.59									
4	93.54	0.40	0.59									
5	92.15	0.20	0.28									
6	90.05	0.65	1.52									
Avg	94.01	0.06	0.17									
BYU 2-year	75.79	0.77	2									

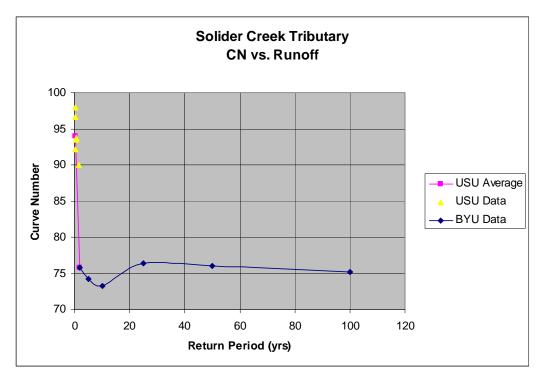


Figure 4-3: Solider Creek Tributary Data Comparison Graph

**Table 4-4: Wattis Branch Comparison Data** 

USU	USU Wattis Branch Data												
Storm	CN	P (in)	T (yrs)										
1	97.63	0.10	0.18										
2	96.88	0.17	0.23										
3	95.44	0.17	0.23										
4	94.43	0.20	0.25										
5	87.95	0.50	0.75										
6	86.42	0.60	1.08										
Avg	93.12	0.29	0.35										
BYU 2-year	75.44	0.01	2										

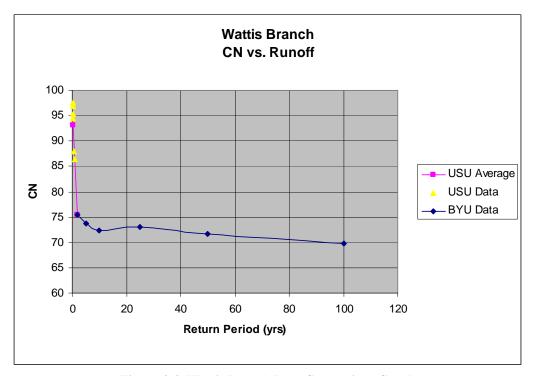


Figure 4-4: Wattis Branch Data Comparison Graph

As discussed in 3.1, the Curve Number generally increases with decreasing return-period (Hawkins 2007). Although there appears to be a hump in the data near the 25-year return-period, the data does generally follow this trend. Using gauged data for storms of a larger size would be a better comparison for this research since the USGS regression equations are only available for 2-year return-period events and greater. Since encountering a storm of larger size is not probable, there is limited data for larger storm events. The USU data however falls where it would be expected on the graph given the small amount of precipitation of each storm. Although it is not an exact fit, the results of this comparison are encouraging since the general shape of the graph is close to what was expected.

## 4.2 Historical Gauged Data CN Calibration Comparison

In an attempt to further confirm the validity of the use of regression equations in CN calibration, CNs calibrated using historical gauged data from Joel William's (Williams 2005) research were compared to results using the regression equations. Three of Williams' forty gauge pairs were selected for comparison. These gauge pairs were considered a good choice for the comparison since the historical precipitation and stream gauge data were well correlated and resulted in reasonable calibrations. Some CN calibrations in Williams' research produced very unreasonable results due to the distance between the rain and stream gauge. To avoid this problem, the selected gauges were located within or very close to the watershed. The gauge pairs include Beaver River near Beaver, Utah, Centerville Creek near Centerville, Utah, and Coal Creek near Cedar City, Utah. The Coal Creek used for this comparison is at a different location than the Coal

Creek Tributary used in the USU research. The location of these watersheds and their gauges can be seen in Figure 4-5 below.

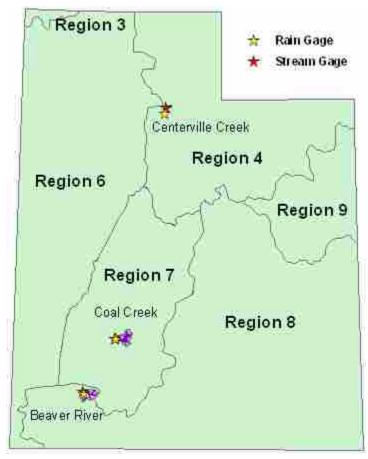


Figure 4-5: Location of Watersheds with Historical Data

The Beaver River watershed gauge coordinates and storm data are shown below in Table 4-5. The hydrograph in Figure 4-6 was developed as a part of Williams' research. This hydrograph depicts the actual storm event rather than an average hydrograph which is an estimation used for design purposes only. The rain gauge (indicated by the yellow star in Figure 4-7) for this watershed is located within the basin, which greatly improves the chances that the rain and stream gauge have recorded data for the same storm.

**Table 4-5: Beaver River CN Calibration Data Summary** 

Precipitation Gage:	Beaver 4 E, Beaver County, UT										
Location (lat/long: d.m.s):	38 17 0 112 34 0										
Stream Gage:	Beaver River near Beaver, UT										
Location (lat/long: d.m.s):	38 16 50 112 34 3										
Date of Storm:	8/7/1987										
Basin Area (mi^2):	91.72										
Total Precipitation (in):	1.4										
Base Flow (cfs):	46										
Peak Flow (cfs):	77										
Runoff Volume (ff*3):	2772000										
UDOT CN:	61.4										
Gaged CN:	64										

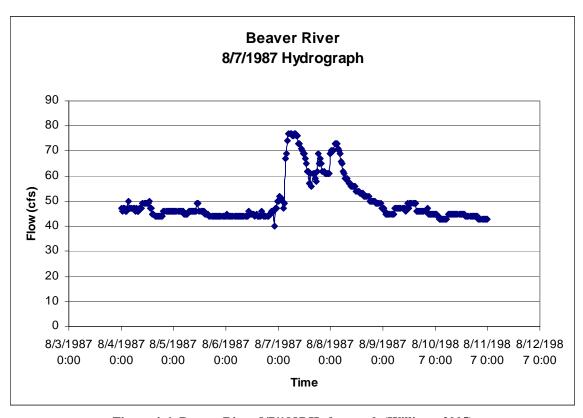


Figure 4-6: Beaver River 8/7/1987 Hydrograph (Williams 2005)

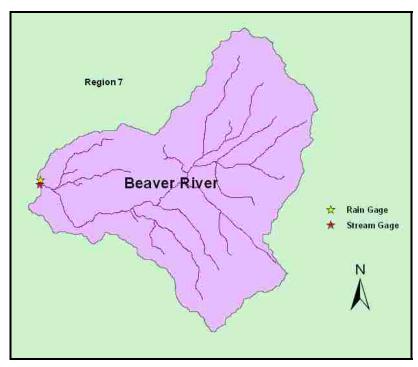


Figure 4-7: Beaver River Gauge Locations

As can be seen in Figure 4-5 and Figure 4-7, Beaver River is located in the NFF Region 7. The watershed was delineated using WMS, and the map files were run through the script using the Region 7 USGS regression equations. The output for Beaver River can be seen in Table 4-6 along with the results for Centerville and Coal Creek. As discussed previously, the precipitation comes from the NOAA Atlas 14 rainfall grid, and the runoff volume is determined using the NFF design hydrograph. The runoff is calculated by dividing the runoff volume by the basin area. With this data, Curve Numbers indexed by return-period could then be back-calculated as seen in Table 4-6.

Table 4-6: Script Output for Historical Gauged Data

Terrenovie:			6		Precipit	alion (m.		Volume of the NFF hydrograph (fft*3)						Attes	Runoff (m.)						
FILENOME	TOOU	CLASS	8 P2 P5 P10 P25 P5				P:50	P#800	V2 V6 V10 V25 V59 V				V408	(min2)	FQ:	R5	RHO	R25	R50	R100	
Beaver River, map	62.2	61.4	1.44	1.76	204	2.44	2.77	3.16	6.E+06	1.E+07	1.E+07	1 E+07	2.E+07	2:E+07	91.72	0.0289	0.0450	0.0579	0.0666	0.0785	0.9919
Centerville Dreek,map	65.9	69.3	1.30	1.59	1.83	2.19	2.50	2.88	8 E+04	1.E+05	2.E485	3.E+05	3,E+05	4.E+05	3,17	0.0105	0.0197	0.0271	0.0368	0.0461	0.8527
Coal Creek map	67.9	81.9	1.48	1.83	2013	2.56	2.92	3.32	6.E+96	1.E+07	2.E+07	3.E+07	3.E+07	4.E407	77.77	0.0343	0.0685	0.0965	0.1429	0.1848	0.2283
Backgard Red ON																					
FILENRALE	_	- 1	SE KORE	FREC !	(4)																
Market Market Control of the Control	-	_ (الآليا _	- Leggle	100	1200	- Later Street															
Beaver River, map	65.99	62.16	5923	54.51	51.60	48.47															
Centerville Creek map	6427	61.90	59.08	55.16	52.28	48.91															
Doal Creek map	65.80	63.22	60.80	67.70	EE 10	53.00															

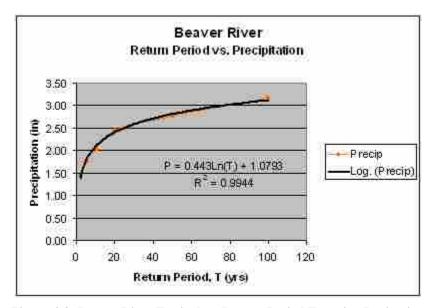


Figure 4-8: Beaver River Equivalent Return-Period Equation Derivation

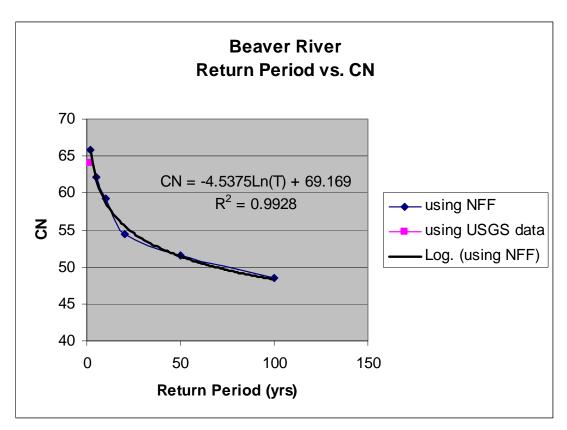


Figure 4-9: Equivalent CN Equation Derivation

In order to determine how the USGS regression equation calibration method compares to calibration using gauged data, the NOAA precipitation data was plotted against the return-period and a trend line was then plotted as seen in Figure 4-8. With the associated equation, the equivalent return-period was calculated for the gauged precipitation. As noted in Table 4-5, the total precipitation for the storm occurring on 8/7/1987 was 1.4 inches. Setting P equal to 1.4 inches in the following equation results in an equivalent return-period of T=2.1 years.

$$P = 0.443Ln(T) + 1.0793 (4-2)$$

The calibrated CNs were then plotted against the return-period and again, a trend line was drawn as seen in Figure 4-9. The trend line equation, Equation 4-3, is shown below.

$$CN = -4.5375Ln(T) + 69.169$$
 (4-3)

The equivalent return-period for the historical data of 2.1 years was then inserted into Equation 4-2 in order to estimate the CN that would be calibrated for the measured precipitation data if calibrated using the regional regression equations. The resulting calibrated CN was 65.9. This estimation is an increase of 2.9% from the CN of 64, calibrated using the historical gauged data.

The same procedure was similarly used for the Centerville Creek and Coal Creek watersheds. The maps, tables, graphs and equations for these watersheds can be viewed in Appendix E. For Centerville Creek the gauged CN calibration was 69 where as the regression equation calibration was 67, a 2.9% decrease. For Coal Creek, the gauged CN was 69.1 and the regression equation CN was 66.9, a 3.3% decrease. With such consistent results, it appeared that CN calibration using the regional regression equations was a viable option where gauged data are unavailable.

## 4.3 CN Calibration using USGS Regression Equations

After determining that using the USGS regression equations was an appropriate means of calibration, roughly ten watersheds from each NFF region were selected for research in order to determine if any regional trends exist and what recommendations

could be made to UDOT for the improvement of their CN table. As mentioned previously, the Utah USGS regression equations were developed for rural regions only, thus the watersheds selected were located in undeveloped areas. Although site visits were not possible to verify that no diversions were present, a topographic map was used to select watersheds with no visible upstream diversions. Only watersheds with a basin area greater than 10 square miles were selected for the study. These selection guidelines were used in order to stay within the limitations of the regression equations (Table 2-3) and the CN method. A map of the selected watersheds can be seen in Figure 4-10 below.

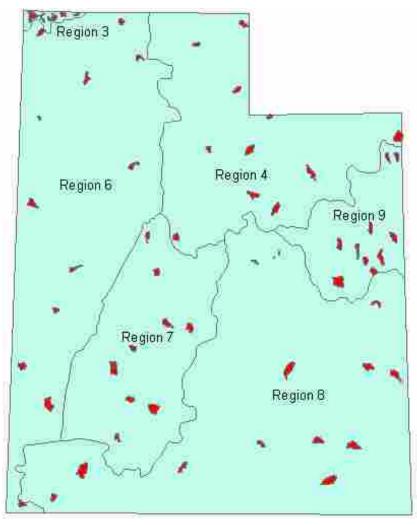


Figure 4-10: Watersheds used in Regional Trend Study

The catalog developed by Dyer (Dyer 2006) was again used for data acquisition and watershed selection. The selected watersheds were processed with the script. The output and back-calculated Curve Numbers for each watershed and return-period have been indexed by region and can be seen in the following section. Region 3 is so small that watersheds wholly contained in the region that were of reasonable size so as to produce reliable results were difficult to find. Therefore, Region 3 was neglected in the study.

## 5 Results

The results for the regional trend study for the state of Utah are shown by NFF region in Table 5-1 through Table 5-5. The data graphs are shown in Figure 5-1 through Figure 5-5. Definitions for the majority of the output were outlined previously in 3.3. In this section, there is a new portion included in the output tables entitled "Scalars". UDOT has requested suggestions on how to modify their CN table to account for the local conditions in various locations in Utah. The values in the scalar table were calculated by taking the back-calculated or calibrated CN of each return-period and dividing by the composite CN that was calculated using the unmodified UDOT CN table for the respective watershed. Near the bottom of the scalar table, there are two rows entitled "WTD AVG" and "AVG". "WTD AVG" indicates a weighted average of all the watershed scalars for a particular return-period. The weighted average was weighted by watershed area. The row entitled "AVG" is a straight average of the calculated scalars for each return-period. Theoretically, if enough watersheds were calibrated using this method in each NFF region, composite CNs calculated using the unmodified UDOT CN table could be multiplied by the weighted average scalar of the NFF region to adjust for the local conditions as well as the return-period for which the project is being designed.

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Table 5-1: Region 4 Calibrated Curve Numbers and Scalars

FILENAME					Precipita	ation (in			V.	olume o	the NFF	hydrog	raph (ft^	3)	Area			Runo	ff (in)		
FILENAME	UDOT	CLASS	P2	P5	P10	P25	P50	P100	V2	V5	V10	V25	V50	V100	(mi^2)	R2	R5	R10	R25	R50	R100
CowHollow.map	68.1	69.3	1.09	1,35	1.58	1.91	2 19	2.51	5.E+05	7 E±05	9.E+05	1.E+06	1.E+06	1.E+06	7.75	0.0264	0.0391	0.0484	0.0591	0.0684	0.0762
CrousseCreek.map	68.1	66.2	0.94	1.18	140	1.72	1.97	2.26	2 E+05	3-E+06	4 E+06	6.E+06	6.E+06	6.E+06	33.79	0.0205	0.0347	0.0458	0.0596	0.0715	0.0823
DeepCreek.map	63.1	62.6	0.94	1.17	1 38	1.67	1.93	2.22	1.E+06	2.E+06	3.E+06	3.E+06	4.E+06	4.E+06	20.62	0.0255	0.0416	0.0539	0.0689	0.0818	0.0932
NorthForkProvoRiver.map	65.2	64.8	1.15	1.42	1 65	2 00	2.29	2.61	2 E+06	3 E+06	4.E÷06	5 E÷06	5 E+06	5 E-06	24.40	0.0426	0.0588	0.0703	0.0832	0.0946	0.104
ProvidenceCanyon.map	812	77.8	1.15	1.41	1.65	2.00	2.29	2.60	4.E+05	6 E+05	8.E+06	1.E+06	1.E+06	1.E+06	10.05	0.0155	0.0257	0.0336	0.0432	0.0514	0.0587
RabbltCreek.map	78.6	68 9	0.62	1.03	1.21	1.10	1.71	1.96	7.E+05	1E+06	2 E+06	2 E+06	3.E+06	3 E+06	13.71	0.0213	0.0396	0.0545	0.0739	0.0906	0.106
SamsCanyon.map	60	59	0.90	1.12	1.31	1.59	1.83	2.12	1.E+06	2 E+06	3.E+06	4 E+06	4 E+06	5.E+06	23.25	0.0231	0.0385	0 0503	0.0648	0.0774	0 0888
SpringCreek.map	76.7	85.6	0.87	1.08	1.27	1.54	177	2 03	6:E+05	1 E+06	2 E+06	2.E+06	3 E+06	3 E+06	13.31	0.0206	0.0370	0.0501	0.0669	0.0813	0.094
WaterHollow.map	78.3	72.7	0.75	0.95	1 13	1.38	1.59	1.85	1.E+06	3 E+06	4 E+06	5.E+06	6.E+06	7.E+06	24.46	0.0252	0.0465	0.0637	0.0862	0 1056	0.1236
WhitePineCanyon.map	65.6	73.7	1.15	1.39	1.61	1.93	2.20	2.51	4 E+06	6 E+05	7 E+05	9:E±05	1.E+06	1E+06	9.90	0.0167	0.0266	0.0320	0.0396	0.0461	0.0516

FILENAME	Ų.	. 8	ackcalc	ulated C	N)	
FILENAME	CN2	CN5	CN10	CN25	CN50	CN100
CowHollow.map	72.47	68.81	65,57	61 15	57.93	54.46
CrousseCreek.map	74 98	71.57	68.51	64.32	61.30	58.02
DeepCreek.map	75.66	72.56	69.76	65.82	62.75	59.35
NorthForkProvoRiver.map	73.21	69.50	66.25	61.75	58.47	55.06
ProvidenceCanyon.map	69 40	65.87	62.74	58.47	55.29	52 10
RabbitCreek.map	78.00	75.52	73.06	69.48	66.80	63.83
SamsCanyon.map	76.44	73.40	70.70	66.87	63.86	60.39
SpringCreek.map	76.68	74 07	71.58	67 93	65.18	62 10
WaterHollow.map	80.25	77.90	75.69	72.36	69.83	66 63
WhitePineCanyon.map	69.67	66.35	63.30	59.05	55.92	52 69

	1 01         0.96         0.90         0.85         0.81           1 05         1.01         0.94         0.90         0.86           1.15         1.11         1.04         0.99         0.9           1 07         1.02         0.95         0.90         0.8           0.81         0.77         0.72         0.68         0.6           0.96         0.93         0.88         0.85         0.8           1.22         1.18         1.11         1.06         1.0           0.97         0.93         0.89         0.85         0.8										
52	S5	S10	S25	550	S100						
1.05	1.01	0.96	0.90	0.85	0.80						
1 10	1.05	1.01	0.94	0.90	0.65						
1.20	1.15	441	1.04	0.99	0.94						
1.12	1.07	1.02	0.95	0.90	0.84						
0.85	0.81	0.77	0.72	0.68	0.64						
0.99	0.96	0.93	0.88	0.85	0.81						
1.27	1 22	1 18	1.11	1.06	1.01						
1:00	0.97	0.93	0.89	0.85	0.81						
1.02	0.99	0.97	0.92	0.89	0.85						
1.06	1.01	0.96	0.90	0.85	0.80						
1.09	1.05	1.01	0.95	0.91	0.86						

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Table 5-2: Region 6 Calibrated Curve Numbers and Scalars

ACTUAL SERVICE CONT.				9	Precipita	ation (in)	N.		V	olume of	the NEE	hydrog	raph (ft	31	Area			Runo	ff (in)		
FILENAME	UDOT	CLASS	P2	P5	P10	P25	P50	P100	V2	V5	V10	V25	V50	V100	(mi*2)	R2	R5	R10	R25	R50	R100
BlackhamCanyon.map	61.4	60.2	0.88	1.11	1.30	1.58	1.81	2.08	-	3.E+05	5.E+05	1.E+06	2 E+06	4.E+06	10.67		0.0108	0.0219	0.0460	0.0799	0.1518
BoxCanyon.map	71	69.6	0.84	1.05	1.23	1.49	1.70	1.94	-	7.E+05	1.E+06	3.E+06	6.E+06	1.E+07	12.30	14000	0.0250	0.0517	0.1115	0 1941	0.3654
BoxElderWash.map	75.6	73.9	1.06	1.28	1.48	1.78	2.00	2 28		6 E+05	1.E+06.	2 E+06	4 E+06	7 E+08	14.72	5 41114	0.0172	0.0326	0.0882	0 1192	0 2166
CaveCreek.map	88.8	87	0.67	0.85	1.00	1.23	1.41	1.63	2222	6 E+64	2 E+05	4 E+05	8.E+05	2 E+06	285	2000000	0.0095	0.0291	0.0672	0 1136	0 2657
FourmileDraw.map	74.3	73.4	0.98	1.23	1.45	1.77	2 02	2 32		2 E+06	4 E+06	8 E+06	1E+07	2 E+07	35.05		0.0253	0.0438	0.0949	0 1687	0.2778
JacksonWash.map	73	73	1 15	1:44	1 69	2.05	2.34	2.66		4 E+05	7 E+05	1E+06	3 E+06	4 E+06	17:44		0.0104	0.0178	0.0358	0.0628	0.1099
KimbellCreek.map	69.4	64.8	0.95	1,19	1.41	1.71	1.96	2 25	37725	5 E+05	9 E+05	2 E+06	3 E+06	6 E+06	13.59	3*****	0.0149	0.0293	0.0622	0 1086	0 2005
NewfoundlandBasin.map	68.8	64.5	0.64	0.81	0.95	1.17	1.36	1.55		1 E+06	3 E+06	7 E±06	1.E+07	2 E+07	15.22		0.0279	0 0744	0:1873	0.3276	0.6384
NorthCanyon.map	65.7	64.4	0.93	1.15	1.34	1.62	1.86	2.13	-	5 E+05	1 E+06	3.E+06	6.E+06	1 E+07	10.45	) <b></b>	0.0217	0.0587	0.1436	0.2492	0.5040
RoseRanchCreek.map	79.9	69	0.83	1.04	1.22	1.48	1.70	1.96		5-E+05	1 E+06	3:E+06	5.E+06	1 E+07	11.29	*****	0.0173	0.0473	0.1174	0.2042	0.4112
FILENAME			ackcalci		4	III-A-III-A-III-A-III-A				200		LAR			b						
A.A. S.	CN2	CN5	CN10	CN25	CN50	CN100															
BlackhamCanyon.map			2000	and the same of		Hiterary Arterior			\$2	S5	S10	\$25	\$50	\$100							
	<del>2000</del>	69.34	67.46	65.28	64.40	64.71				1 13	1.10	1.06	1 05	1.05							
BoxCanyon.map	2000000 	73.09	72:45	65.28 72.12	72.90	64.71 75.43				1 13 1 03	1 10 1 02	1.06	1.05	1.05 1.08							
BoxCanyon.map BoxElderWash.map		69.34 73.09 67.10	67.46 72.45 65.63	and the same of	72.90 64.11	64.71 75.43 64.94				1 13 1.03 0.69	1 10 1 02 0 87	1.06 1.02 0.85	1.05 1.03 0.85	1.05 1.06 0.86							
BoxCanyon.map BoxElderWash.map CaveCreek.map	-	73.09 67.10 75.04	72.45 65.63 74.81	65.28 72.12 64.25 73.93	72.90 64.11 73.63	64.71 76.43 64.94 77.03			+	1 13 1.03 0.89 0.85	1.10 1.02 0.87 0.84	1.06 1.02 0.85 0.83	1 05 1 03 0 85 0 83	1.05 1.06 0.86 0.87							
BoxCanyon.map BoxElderWash.map	****	73.09	72:45	65.28 72.12	72.90 64.11	64.71 75.43 64.94				1 13 1 03 0 89 0 86 0 93	1 10 1 02 0 87 0 84 0 91	1.06 1.02 0.85 0.83 0.89	1 05 1 03 0 85 0 83 0 89	1.05 1.08 0.86 0.87 0.90							
BoxCanyon.map BoxElderWash.map CaveCreek.map FourmileDraw.map JacksonWash.map		73.09 67.10 75.04 69.28 62.76	72.45 65.63 74.81 67.39 59.96	65.28 72.12 64.25 73.93 66.22 57.01	72.90 64.11 73.63 66.39 55.56	64.71 76.43 64.94 77.03 66.78 54.74			70000 70000 70000	1 13 1 03 0 89 0 85 0 93 0 86	1 10 1 02 0 87 0 84 0 91 0 82	1.06 1.02 0.85 0.83 0.89 0.78	1 05 1 03 0 85 0 83 0 89 0 76	1.05 1.08 0.86 0.87 0.90 0.75							
BoxCanyon.map BoxElderWash.map CaveCreek.map FourmileDraw.map JacksonWash.map KimbellCreek.map		73.09 67.10 75.04	72.45 65.63 74.81	65.28 72.12 64.25 73.93	72.90 64.11 73.63	64.71 76.43 64.94 77.03 66.78 54.74 64.60			2000 2000 2000 2000 2000 2000	1 13 1 03 0 89 0 86 0 93 0 86 0 99	1 10 1 02 0 87 0 84 0 91	1.06 1.02 0.85 0.83 0.89	1 05 1 03 0 85 0 83 0 89 0 76 0 92	1.05 1.06 0.86 0.87 0.90 0.75 0.93							
BoxCanyon.map BoxElderWash.map CaveCreek.map FourmileDraw.map JacksonWash.map		73.09 67.10 75.04 69.28 62.76	72.45 65.63 74.81 67.39 59.96	65.28 72.12 64.25 73.93 66.22 57.01	72.90 64.11 73.63 66.39 55.56	64.71 76.43 64.94 77.03 66.78 54.74 64.60 88.31				1 13 1 03 0 89 0 85 0 93 0 86	1 10 1 02 0 87 0 84 0 91 0 82 0 96 1 17	1.06 1.02 0.85 0.83 0.89 0.78	1 05 1 03 0 85 0 83 0 89 0 76	1.05 1.08 0.86 0.87 0.90 0.75							
BoxCanyon.map BoxElderWash.map CaveCreek.map FourmileDraw.map JacksonWash.map KimbellCreek.map	2000 2000 2000 2000 2000 2000	73.09 67.10 75.04 69.28 62.76	72.45 65.63 74.81 67.39 59.96	65.28 72.12 64.25 73.93 66.22 57.01 64.59	72.90 64.11 73.63 66.39 55.56 64.02	64.71 76.43 64.94 77.03 66.78 54.74 64.60				1 13 1 03 0 89 0 86 0 93 0 86 0 99	1 10 1 02 0 87 0 84 0 91 0 82 0 96	1.06 1.02 0.85 0.83 0.89 0.78 0.93	1 05 1 03 0 85 0 83 0 89 0 76 0 92	1.05 1.06 0.86 0.87 0.90 0.75 0.93							

WTD AVG: AVG:

Table 5-3: Region 7 Calibrated Curve Numbers and Scalars

THE ENLANCE:					Precipit	ition (in			V	olume of	the NFF	hydrog	raph (ft	3)	Area			Runo	ff (in)		
FILENAME	UDOT	CLASS	P2	P5	P10	P25	P50	P100	V2	V5	V10	V25	V知	V100	(mI^2)	R2	R5	R10	R25	R50	R100
CunninghamWash.map	75.1	71.1	1.01	1.26	1.47	1.78	2.03	2 33	3.E+05	8 E+05	1.E+06	3.E+06	3 E+06	4 E+06	19:57	0.007	0.017	0.027	0.055	0.071	0.090
DogValleyCreek.map	69.6	66.5	0.95	1.19	1 38	1.67	1.90	2.17	2.E+06	6 E+05	8.E÷05	2.E+06	2.E+06	3 E+06	11.95	0.007	0.018	0.029	0.063	0.064	0 108
FoolCreek.map	60	59	0.99	1.23	1.43	1.70	1.93	2.21	2.E+05	5.E+05	8.E+05	2.E+06	2.E+06	3.E+06	14.97	0.006	0.015	0.023	0.053	0.069	0.088
HorseCreek.map	62.3	62.6	1.42	1.76	2.05	2.47	2.83	3.23	6:E+05	1 E+08	2 E+06	2 E+06	3.E+06	4 E+06	10.88	0.025	0.049	0.071	0.092	0.122	0.158
JerichoWash.map	79.5	78.6	0.83	1.03	1.20	1.45	1 65	1.89	2 E+05	6 E+05	1.E+06.	3.E+06	4 E+06	6.E+08	10.17	0.008	0.024	0.043	0 140	0.187	0.243
LeesSpringWash.map	65.8	64.3	1.00	1.24	1.45	1.75	2.00	2 29	3 E-05	8 E+06	1 E-06	3.E+06	3.E+06	4 E+06	15.56	0.009	0.021	0.034	0.070	0.092	0.117
LostCreek.map	75.2	74.4	1 04	1.29	1.51	1 83	2.09	2.42	9.E+05	2 E+06	3.E+06	4 E+06	5 E+06	6 E+06	34 90	0.011	0.022	0.032	0.050	0.062	0.077
RidgeCreek.map	78.6	78.4	1.06	1.30	1.52	1 82	2.07	2.38	7.E+05	1.E+06	2 E+06	4 E+06	5.E+06	6.E+06	19.69	0.014	0.031	0.047	0.084	0.108	0.136
SoliderCreek.map	74.4	71.7	0.84	1.04	1.22	1.47	1 69	1.95	3.E+05	8 E+05	1 E+06	3.E+06	4 E+06	5.E+06	20 01	0.007	0.016	0.029	0.070	0.091	0.114
WaterCanyon.map	75.9	71.0	0.97	1.22	1.42	1.72	1.96	2.25	2E-05	5 E±05	8:E+05	2.E+06	2 E+06	3.E+06	13.77	0.006	0.015	0.025	0.056	0.073	0.094

FILENAME		В	ackcalc	ulated C	N.	· ·
FILENAME	CN2	CN5	CN10	CN25	CN50	CN100
CunninghamWash.map	70.61	67 40	64.95	62.90	60.46	57.63
DogValleyCreek.map	71.86	69.01	66.84	65.40	63.26	60.89
FoolCreek.map	70.54	67.61	65.35	63.93	61.68	59 14
HorseCreek.map	65,66	62.63	60.08	55.92	53.56	51 10
JerichoWash.map	74.99	73.42	72.14	74.61	73.36	71.81
LeesSpringWash.map	71.40	68.54	66.33	64 62	62 34	59 79
LostCreek.map	71.06	67.70	65.01	61.64	58.83	55 61
RidgeCreek.map	71.21	68.71	66 59	64 60	62.32	59 73
SoliderCreek.map	74.64	72 13	70.06	69.43	67.16	64 51
WaterCanyon.map	71.04	67 90	65.53	63.86	61.57	58.87

			SCA	LAR		
	S2	S5	S10	\$25	\$50	S100
	0.94	0.90	0.86	0.84	0.61	0.77
	1.03	0.99	0.96	0.94	0.91	0.88
	1.18	1 13	1.09	1.07	1 03	0.99
	1.05	1.00	0.96	0.90	0.86	0.82
	0.94	0.92	0.91	0.94	0.92	0.90
Ī	1.09	1.04	1.01	0.98	0.95	0.91
	0.94	0.90	0.86	0.82	0.78	0.74
	0.91	0.87	0.85	0.82	0.79	0.76
	1.00	0.97	0.94	0.93	0.90	0.87
	0.94	0.89	0.86	0.84	0.81	0.78
VTD AVG:	0.99	0.95	0.92	0.89	0.86	0.82
AVG:	1.00	0.96	0.93	0.91	0.88	0.84

7

**Table 5-4: Region 8 Calibrated Curve Numbers and Scalars** 

FILENAME				. 1	Precipit	ation (in			V	olume o	the NF	hydrog	raph (ft	3)	Area			Runo	ff (in)		
FILENAME	UDOT	CLASS	P2	P5	P10	P25	P50	P100	-V2	V5	V10	V25	V知	V100	(mI^2)	R2	R5	R10	R25	R50	R100
ClayCanyon.map	74.7	55.3	0.70	0.89	1.06	1.32	1.53	1.81	3 E+06	9 E+06	1 E+07	2 E+07	3 E+07	4 E+07	11.78	0.1237	0.3135	0.4935	0.7994	1.0884	1,4161
CettonwoodWash.map	67.2	64.8	0.87	1:11	1.31	1.58	1.83	2.08	4 E+06	1 E±07	2 E+07	3:E±07	4.E+07	6 E+07	13.20	0.1345	0.3787	0 6259	1.0511	1.4598	1 9376
DryCanyon.map	78.8	73.1	0.75	0.95	1.12	1.38	1.60	1.87	4.E÷06	9.E+06	1.E+07	2.E+07	3.E+07	4.E+07	10.57	0.1461	0.3647	0.5702	0.9193	1.2484	1.6201
HammondCanyon.map	75.8	74.5	0.98	1.23	1.46	1.78	2.06	2.37	5 E+06	1 E+07	2 E+07	2 E+07	3.E+07	4 E=07	27 07	0.0755	0.1672	0.2473	0.3809	0.5040	0.6373
Hanksville.map	66.9	64.6	0.67	0.86	1.02	1.27	1 48	1.76	1.E+07	3 E+07	4.E+07	6.E+07	8 E+07	1.E+08		0.1315	0.3118	0.4759	0.7500	1.0046	1.2860
HunterCanyon.map	87.8	81.9	0.75	0.95	1 13	1.40	1.54	1.92	5 E+06	1 E+07	2 E+07	3 E+07	4 E+07	6 E+07	22 23	0.1031	0.2508	0.3875	0.6174	0.8324	1.0726
LaSal.map	51.5	52.3	1 03	1.28	1.50	1.84	2 12	2.45	3.E+06	6 E+06	9.E+06	1 E+07	2 E+07	2 E+07	23 29	0.0546	0.1150	0.1661	0.2518	0.3301	0.4136
LongCanyon.map	87.2	63 6	0.80	1.01	1.20	1.46	1.72	2.01	5 E+06	1.E+07	2 E+07	3.E+07	4 E+67	5.E+07	21.17	0.0962	0 2289	0.3499	0.5532	0.7426	0.9528
RoundValleyDraw.map	82.3	79.1	0.89	1,13	1.33	1.52	1.87	2.17		9 E+06					16.20		0.2378		0.5657	0.7570	0.9681
SlickhornCanyon.map	72.6	711	0.81	1.03	1.23	1.51	1,76	2.04	9.E+06	2 E±07	3.E+07.	5.E+07	7 E±07	8.E+07	38.07	0.1059	0.2423	0.3638	0.5662	0.7532	0.9575
									Ü	1											
FILENAME		В	ackcalc	ufated C	N.						SCA	LAR			Á						
FILLIAMIC	CN2	CN5	CN10	CN25	CN50	CH100			S2	S5	S10	S25	\$50	S100							
ClayCanyon.map	89 18	91.70	92.87	94.37	95.54	96.32			1 19	1_23	1.24	1.26	1.28	1 29							
CottonwoodWash.map	85.90	89.64	91,75	94.42	96.57	98.77			1.28	1.33	1.37	141	1.44	1.47	į						
DryCanyon.map	89.07	92.04	93.43	95.21	96:62	97.71	j l		1.13	1 17	1 19	1.21	1.23	1.24							
HammondCanyon.map	79.84	80.03	79.34	78.41	77.66	76.45			1.05	1.06	1.05	1.03	1.02	1.01							
Hanksville.map	90 16	92.30	93.14	94.28	95.13	95.36	ľ		1.35	1.38	1.39	1.41	1.42	1.43							
HunterCanyon.map	86.65	88.79	89.47	90.17	90.67	90.73			0.99	1.01	1.02	1.03	1.03	1.03							
LaSal.map	76.96	76.26	74.91	73.02	71.52	69.43			1.49	1 48	1.45	1.42	1.39	1.35							
LongCanyon.map	85.42	86.91	87.27	87 68	87.97	87 70			1:27	1 29	1 30	1.30	1.31	1.31							
RoundValleyDraw.map	83.61	85.08	85.43	85 80	86.10	85.80			1.02	1 03	1.04	1:04	1.05	1.04							
SlickhornCanyon.map	86 67	87 02	87 24	87.61	87.69	87 40			1 18	1 20	1.20	1.21	1.21	1.20							
								UNIC TE CIEFE							į.						
							W	TD AVG:	1,000	1.23	1.23	1.23	1.24	1.23							
								AVG:	1.20	1.22	1,22	1.23	1.24	1.24							

 $\propto$ 

Table 5-5: Region 9 Calibrated Numbers and Scalars

FILENAME					Precipita	ation (in			V/e	olume of	the NFF	hydrog	raph (ft/	13)	Area			Runo	ff (in)		
FILENAME	UDOT	CLASS	P2	P5	P10	P25	P50	P100	V2	V5	V10	V25	V50	V100	(mi^2)	R2	R5	R10	R25	R50	R100
AgencyWash.map	64	61.8	0.77	0.97	1.14	1.40	1.62	1.89	8.E+05	2 E±06	3.E+06:	6.E+06	1.E+07	1.E+07	13:14	0.028	0.063	0.103	0 191	0.317	0.369
lorenceCreek.map	54.5	64	0.86	1.09	1.28	1.67	1.62	2.14	2 E+06	5 E+06	7.E+06	1.E±07	1.E±07	2.E+07	40.50	0.025	0.048	0.069	0.107	0.161	0.175
ongWash.map	62.3	60.9	0.75	0.95	1_12	1.38	1.59	1.84	1.E+06	2.E+06	4 E+06	7.E+06	1.E+07	1.E+07	19.98	0.022	0.048	0.079	0.145	0.242	0.281
MainBottomCanyon.map	62.8	61.2	0.81	1.01	1.20	1.47	1.70	1.98	6 E+05	1 E+06	2 E+06	4.E÷06	6 E+06	7 E+06	13 28	0.021	0.045	0.071	0 121	0 185	0.215
MeadowCreek.map	77.5	73.3	0.91	1,13	1.33	1.63	1.88	2.18	7.E+05	1 E+06	2 E+06	3 E+06	5.E+06	6.E+06	15.26	0.020	0.042	0.062	0.098	0.139	0.161
RedWash.map	77-1	74	0.77	0.98	1.17	1.44	1 65	1.91	3.E+05	8 E+05	1 E+06	3.E+06	5.E+06	5.E+06	11 09	0.013	0.031	0.055	0.112	0.205	0.236
SandCreek.map	66	64	0.69	0.88	1.04	1.28	1.47	1.72	7.E+05	2 E+06	3.E+06	6 E+06	1.E+07	1 E+07	16,25	0.017	0 042	0.074	0 156	0.295	0.343
SeepCanyon.map	75.6	73.8	0.84	1.06	1.24	1.52	1.75	2 04	2E-06	4 E+06	6 E+06	9 E+06	1 E+07	2 E+07	22 05	0.035	0.073	0.110	0 180	0.268	0.310
StonebridgeDraw.map	63.3	61.6	0.78	1:00	1 18	1.46	1:68	1.93	4.E+05	9 E±05	1.E+06	3.E+06	5 E+06	6.E+06	10.92	0.014	0.034	0.058	0 112	0.194	0.226
abyagaCanyon.map	65.6	62.8	0.74	0.93	1 10	1.35	1:57	1.83	7.E+05	2 E+06	3 E+06	6:E±06	9 E+06	1E+07	18.24	0.017	0.039	0.056	0 127	0.222	0.257

FILENAME	Ų,	В	ackcalc	ulated C	N)	
FILENAME	CN2	CN5	CN10	CN25	CN50	CN100
AgencyWash.map	80.30	79.19	78.35	77.96	79.02	76.31
FlorenceCreek.map	77.50	75:11	73.00	70.34	68.64	66.01
LongWash.map	79.79	78.21	77.04	76.19	76.84	74.07
MainBottomCanyon.map	78.28	76.45	74.85	73 07	72.47	69.27
MeadowCreek.map	75 68	73.46	71.38	68 71	67/09	63.71
RedWash.map	77 69	75.53	74.04	73.08	74.07	71 33
SandCreek.map	80.67	79 24	78.48	78.66	80.77	78.29
SeepCanyon.map	79.31	77.96	76.67	75.28	75.01	72 10
StonebridgeDraw.map	77.71	75.56	73.98	72.73	73.23	70.47
TabyagaCanyon.map	79.38	77 58	76.38	75:57	76.28	73 28

			SCA	LAR		_
	52	S5	S10	S25	S50	S100
Ī	1.25	1.24	1 22	1.22	1.23	1 19
ı	1.20	1 16	1.13	1:09	1.06	1.01
I	1.28	1.26	1.24	1.22	1.23	1.19
	1.25	1.22	1.19	1.16	1.15	1.10
l	0.98	0.95	0.92	0.89	0.87	0.82
į	1.01	0.98	0.96	0.95	0.96	0.93
	1 22	1 20	1 19	1.19	1.22	1.19
	1.05	1 03	1.01	1:00	0.99	0.95
Ī	1.23	1 19	1 17	1.15	1.16	1.11
	1.21	1 13	1.16	1.15	1 16	1.11
/TD AVG:	1.17	1.14	1.12	1.10	1.10	1.05
AVG	1 17	4 44	1.12	1.10	1.10	1.06

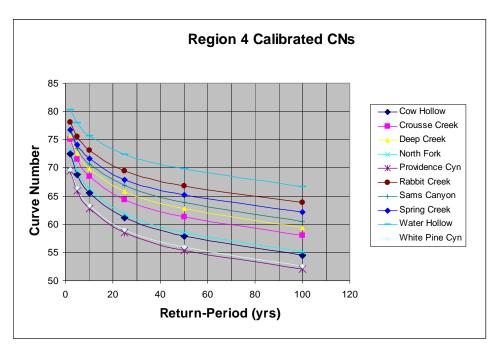


Figure 5-1: Region 4 Calibrated Curve Number Graph

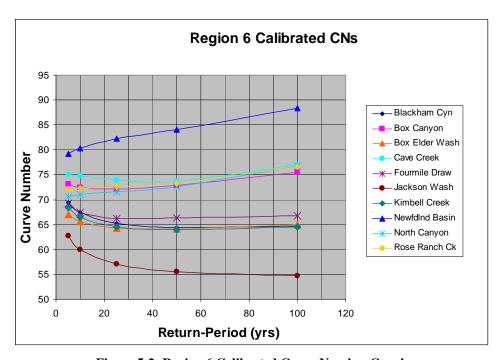


Figure 5-2: Region 6 Calibrated Curve Number Graph

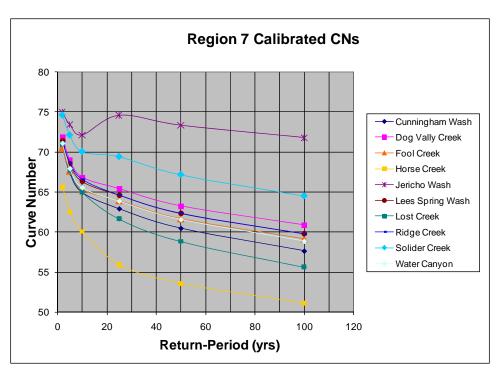


Figure 5-3: Region 7 Calibrated Curve Numbers Graph

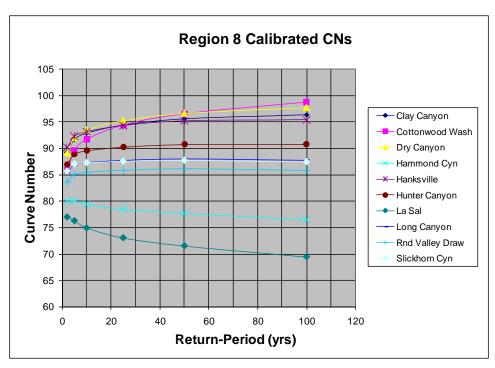


Figure 5-4: Region 8 Calibrated Curve Number Graph

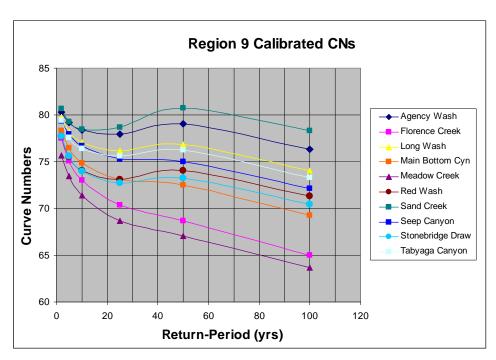


Figure 5-5: Region 9 Calibrated Curve Number Graph

### 6 Discussion

### **6.1** Regional Trends

There are several interesting trends in the calibration results. One of the more obvious trends typical of the majority of the regions is the decrease in the back-calculated Curve Number with increasing return-period. As discussed previously, this trend is to be expected since the greater the return-period, the more precipitation (P) and runoff (R). If both P and R increase with increasing return-period, this would cause the denominator of the CN back-calculation equation as seen below to increase, thus decreasing the CN with increasing return-period.

$$CN = \frac{200}{P + 2R - \sqrt{5PR + 4R^2} + 2} \tag{2-7}$$

The results of Region 4 (see Figure 5-1) are precisely what would be anticipated in all regions. This trend was present in all regions but was not, however, completely consistent in all cases. The calibrated CNs for the majority of the study watersheds in Region 8 and some in Region 6 actually increased with increasing return-period as seen in Figure 5-2 and Figure 5-4 which suggests that the regression equations are predicting a disproportionate amount of runoff for the amount of rainfall present, and that there is a

progressively greater proportion of runoff to rainfall with increasing return-period. It should be noted in Figure 2-11 that both Region 6 and Region 8 are the largest of the Utah NFF Regions and perhaps the least populated. Region 6 extends from the southern to the northern border of western Utah. The climate and geology varies greatly from north to the south in Utah. The consistency of the regression equations over such large areas may be questionable when considering the results of this study.

Upon further investigation, it was noticed that abnormalities in the data appeared to be location based. Watersheds resulting in increasing Curve Numbers with increasing return-period were often clustered together. The Region 6 graph was color-coded in Figure 6-1 to illustrate this concept. The yellow data are those that were considered abnormal. The pink data are those that conformed to the expected trend.

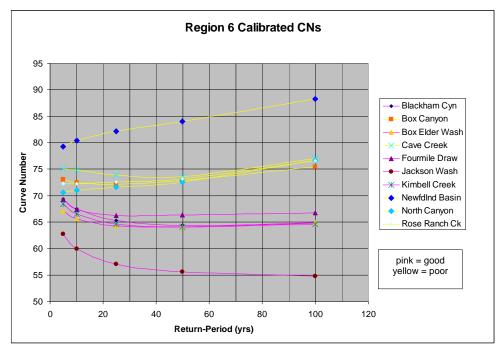


Figure 6-1: Region 6 Data Analysis

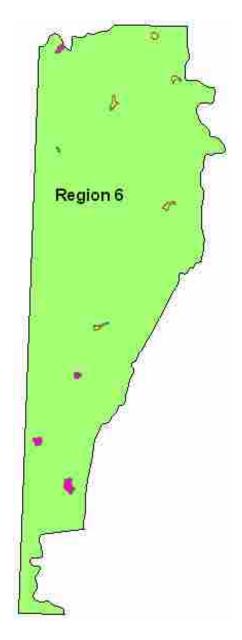


Figure 6-2: Region 6 Watershed Map

All the watersheds processed for Region 6 are shown in Figure 6-1. The watersheds were also color-coded and correlate with the graph in Figure 6-2. Again, the yellow watersheds indicate those with atypical data, and the pink indicate those with normal results. As seen in Figure 6-2, watersheds with results that were as expected seem

to be clustered mostly at the southern end of the region. A single watershed near the top of the region also produced expected results.

Given the size of Region 6, more watersheds would need to be investigated before any concrete conclusions can be drawn. From the available data, however, it appears that there is some correlation between watershed location, and the output data. This correlation could most likely be attributed to the "goodness of fit" of the USGS regression equations. There is only one regression equation per return-period that is applied to the greater part of the western half of the state of Utah. It is quite possible that the regression equations do not fit as well in some areas of the region.

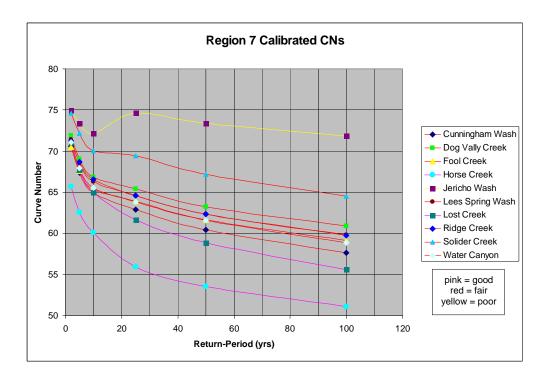


Figure 6-3: Region 7 Data Analysis

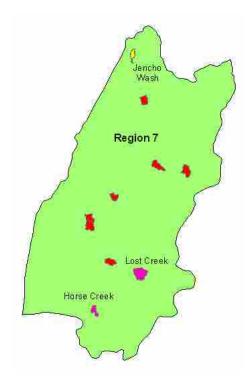


Figure 6-4: Region 7 Watershed Map

Region 7 and Region 9 also contained some data abnormalities. Region 7 in general produced expected results. There is some inclination at the 25-year return-period for the Curve Number to jump up and then continue decreasing with increasing return-period. The more abnormal results came from the "Jericho Wash" watershed as seen in yellow in Figure 6-3 above. The most reasonable results came from the "Lost Creek" and "Horse Creek" watersheds. The relative locations of these watersheds are shown in Figure 6-4. Again it seems that the abnormalities in the data are location based. The best fitting data occurred in the watersheds at the southern end of the region and the less typical data occurred in the farthest north watershed. The clusters of good fitting data also appeared in Region 9 as seen in Figure 6-5 and Figure 6-6.

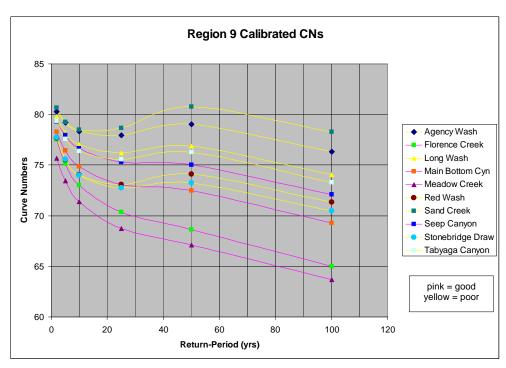


Figure 6-5: Region 9 Data Analysis

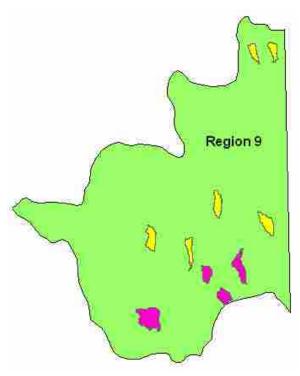


Figure 6-6: Region 9 Watershed Map

Similar to the Region 7 results, the data in Region 9 tends to jump up at the 50-year return-period and then continue decreasing with increasing return-period. This jump in the data could either be due to the regression equations predicting too much runoff or the NOAA rainfall grids predicting too little rainfall at the 50-year return-period and beyond.

To further explore the unexpected data of Region 8, the CN calibrations for the watersheds in Region 8 were also processed using the neighboring Region 7 equations. The Region 7 equations were considered appropriate for use within Region 8 based on observations in past studies (Nelson 2008). When using Region 7 equations the typical trend of decreasing CN with increasing return-period as seen in Figure 6-7 was more prominent. The return of this trend suggests that Region 7 regression equations may be more applicable in some areas of Region 8 than Region 8 equations.

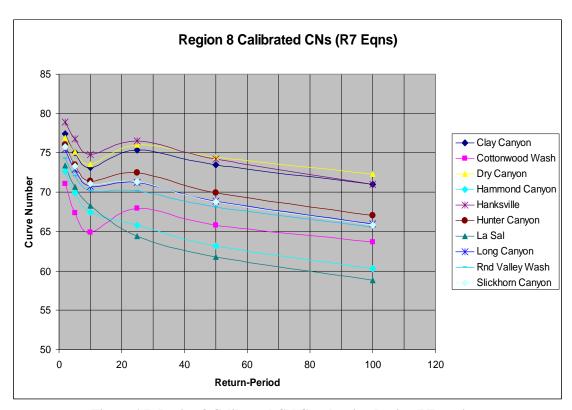


Figure 6-7: Region 8 Calibrated CN Graph using Region 7 Equations

The Presence of regional trends as well as trends within regions was quite strong. The Region 4 results were as expected, which was not surprising since Region 4 is a more populated area of the state. More abundant and better historical rain and stream gauge data would be expected in these areas and would help to develop better interpolations and regression equations. Results for Region 7 and 9 typically held to the anticipated trend of decreasing CN with increasing return-period, but contained an unanticipated hump in the data. Region 8 results and some of the Region 6 results demonstrated a trend completely opposite of what was anticipated. The CNs typically increased with increasing return-period in Region 8. Region 6 and Region 8 are so large that one set of regression equations may not be adequate to sufficiently characterize the runoff in every location for the entire region.

The regression equations that were used in this study are not the most recent generation of regression equations. As mentioned in 2.2.4, in October 2007, the USGS introduced a new set of improved regression equations. Given the variety of results obtained in this study, application of the more recent regression equations is suggested.

#### **6.2** Calibration for UDOT CN Table

Research of CN calibration using the USGS regression equations began with a request from UDOT for a consistent, return-period based CN table that is appropriate for application in the state of Utah. With the research completed thus far some suggestions can be made for the improvement of the UDOT CN table. Table 6-1 is a summary of the weighted average scalars developed in Chapter 5.

Table 6-1: Scalars for UDOT CN Table Calibration

REGION		WEIG	HTED AVE	ERAGE SC	ALAR	
REGION	2YR	5YR	10YR	25YR	50YR	100YR
4	1.09	1.05	1.01	0.95	0.91	0.86
6		0.98	0.96	0.95	0.95	0.97
7	0.99	0.95	0.92	0.89	0.86	0.82
8	1.07	1.04	1.01	1.01	0.97	0.93
9	1.17	1.14	1.12	1.10	1.10	1.05

This table contains the multipliers that UDOT could use to transform composite CNs derived from the use of their CN table to the USGS regression equation calibrated CN for each return-period. These scalars were calculated for each watershed at each return-period by dividing the calibrated CN by the composite CN that was calculated using the UDOT CN table. The values in Table 6-1 are all fairly close to 1 which indicates that the CN values currently used by UDOT are fairly close to those calibrated using the USGS equations. After reviewing this table some regional trends become obvious. The UDOT composite CNs in Region 4 appear to be appropriate for the 10-year return-period (i.e. the scalar is close to 1) but should be adjusted when designing for a higher or lower return-period as indicated in the table. In Regions 6 and 7, the UDOT CNs were consistently high on average when compared to the regression equation calibrated CNs. The UDOT CNs in Regions 8 and 9 were consistently low on average. For improved scalars, a greater number of watersheds should be included in the weighted average and should be calibrated using the more recent regression equations.

### 7 Conclusions

CN calibration using the USGS regression equations were quite consistent when compared to calibrations obtained using measured data. Further comparisons to calibrations using gauged data should be pursued to ensure the accuracy of this method in all regions. The use of this method in large NFF regions should also be further investigated to ensure consistent results.

Although the CNs calibrated through the use of the USGS regression equations inherit the weaknesses of assumptions and estimations made in the development of the regression equations and design hydrograph, the use of USGS regression equations is generally an appropriate method of CN calibration for non-urban areas. This method of CN calibration would be especially useful in cases where calibration is needed on a large scale or when spatially correlated gauged data are unavailable. It is simple, consistent in smaller NFF regions and uses design principles that are common and familiar in industry. Use of the USGS regression equations in calibration provides CNs that are not only calibrated for local application, but are also indexed by return-period, which would be useful for design purposes. While this study was completed for just the state of Utah, the method of CN calibration using the USGS regression equations could be applied in any region in the United States.

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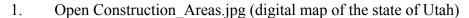
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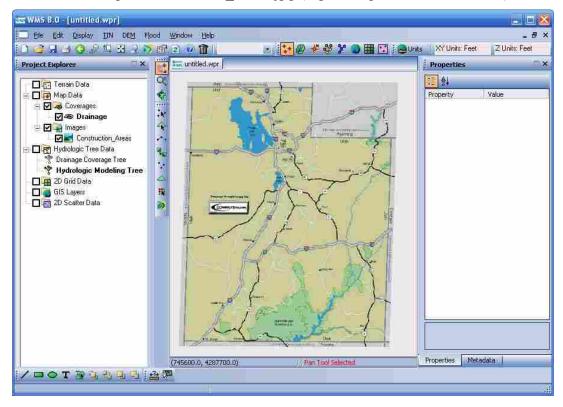
# Appendix A. CN Table Currently used by UDOT

Table A-1: UDOT CN Table (Dyer 2006)

		Soil Type			
LUCODE	Description	A	В	С	D
11	Residential	61	75	83	87
12	Commercial Services	89	92	94	95
13	Industrial	81	88	91	93
14	Transportation Communication	98	98	98	98
16	Mixed Urban or Build-Up Land	75	85	88	98
21	Cropland and Pasture	72	81	88	91
22	Orchards Groves Vineyards Nurseries	62	73	80	85
31	Herbaceous Rangeland	39	61	84	89
32	Shrub and Grass Rangeland	45	66	86	90
33	Mixed Rangeland	72	79	86	92
34	Sagebrush with understory	45	51	68	78
35	Desert Shrubs	50	68	80	86
41	Deciduous Forest Land - Oak and Aspen (80%)	25	32	42	52
42	Evergreen Forest Land	36	60	73	79
43	Mixed Forest Land	36	60	73	79
44	Pinion - Juniper	45	53	75	80
51	Streams and Canals	0	0	0	0
52	Lakes	0	0	0	0
53	Reservoirs	0	0	0	0
54	Bays and Estuaries	0	0	0	0
61	Forested Wetlands	30	55	70	77
62	Nonforested Wetlands	30	58	71	78
71	Dry Salt Flats	74	80	90	92
72	Beaches	50	50	50	50
73	Sandy Areas and Other Beaches	63	77	85	88
74	Bare Exposed Rocks	98	98	98	98
75	Strip Mines, Quarries, and Gravel PIts	77	86	91	94
77	Mixed Barren Lands	77	86	91	94
86	Mixed Rocky – Sparse Junipers	78	87	95	98
87	High Planes	65	69	73	77
91	Perennial Snowfields	0	0	0	0
92	'Glaciers"	0	0	0	0

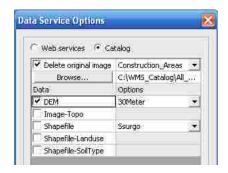
## **Appendix B.** Outlined Calibration Process in WMS





- 2. Choose your watershed and outlet location (coordinates for student projects are found in the "Station Pairs Worksheet.xls")
- 3. Use the "Get Data Tool"
  - a. Drag a box around the watershed to acquire DEM
  - b. When the box pops up click the Catalog option

- c. Browse for the Catalog
- d. Check the DEM box



4. Click "DRAINAGE MODULE TOOL"

DEM>Compute Topaz (choose units, click ok)

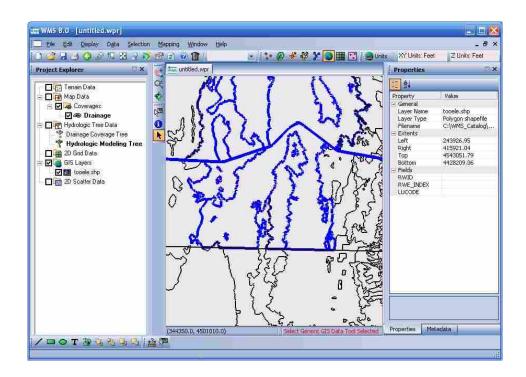
- 5. Create an Outlet at the correct coordinates
- DEM>Delineate Basins Wizard (Click Ok, choose Consistent Units, Click
   Ok)
- 7. Optional: DEM>Delete Null Basin Cells Data
- 8. Right Click on "Coverages" Create New Coverage

Make the coverage a "Land Use" coverage



9. Making sure the new land use coverage is highlighted use the "Get Data Tool"

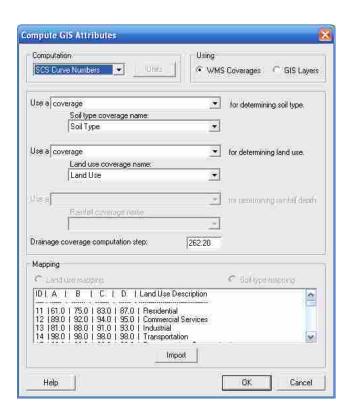
- a. Drag a box around the watershed to acquire Land use
- b. When the box pops up click the Catalog option
- c. Browse for the Catalog
- d. Check the Shapefile- landuse box
- 10. Using the GIS MODULE TOOLS click the "Select Shapes Tool"
- 11. Drag a box around the watershed (some shapes should turn blue)



- 12. Mapping>shapes->feature objects. Next, Next, Finish
- 13. Repeat STEPS 8-12 creating a "SOIL TYPE COVERAGE"
- 14. Using the "HYDROLIC MODELING MODULE"
  - a. Calculators>Compute GIS attribute

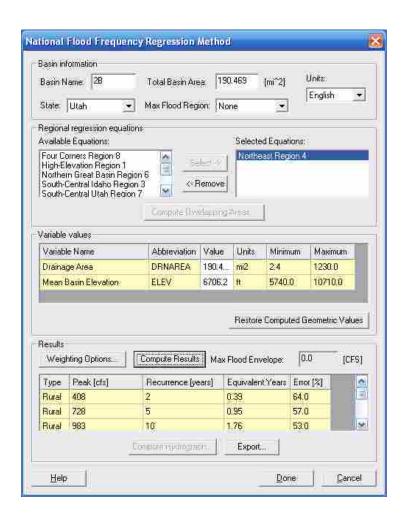
### b. SCS CN

i. Import Class Mapping table click OK



- ii. Record CN Number
- iii. Repeat with UDOT mapping table
- iv. Record CN number
- 15. Right Click on "Coverages" Create New Coverage
  - a. Make the coverage a "NFF" coverage
  - b. Import the NFF regions map
- 16. Convert coordinates (geographic NAD 83 > UTM NAD 83)

17. Using the "HYDROLIC MODELING MODULE" Change the model type in the drop down box to "NFF"

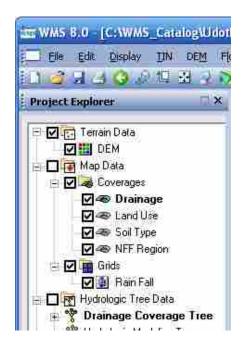


#### 18. Double click on the Basin

- a. Make sure all the information was imported to your NFF modeli.e. the NFF region and the basin areas etc.
- b. Compute results (record results)
- c. Compute hydrographs (record)

Use the Denver Method

19. Open Rainfall depth grid (make sure there extension is .grd so the Program gives you the option to import as a "rainfall depth grid"



Notice the four coverages and the rainfall depth grid

- a. Open HEC-1 script>precipitation to see the calculated value
- b. Divide the computed number by 1000 to obtain the rain fall depth in inches.
- c. Record data in spread sheet
- d. Results will automatically be calculated for CN

(Dyer 2006)

# **Appendix C.** Class Average CN Table

**Table C-1: Class Average CN Table** 

	Class Average CN Table				
LUCODE	Description		Soil	Туре	ļ
LOCODL	Description	Α	В	С	D
11	Residential	60	74	82	87
12	Commercial Services	89	92	94	95
13	Industrial	81	88	91	93
14	Transportation and Communication	76	85	89	91
16	Mixed urban or built up land	77	85	90	93
17	Other urban or built up land	71	82	88	90
21	Cropland and Pasture	49	68	78	84
22	Orchards, Groves, Vineyards, Nurseries	47	67	77	83
23	Confined Feeding Operations	55	63	66	68
24	Other Agricultural Land	62	74	82	86
31	Herbaceous Rangeland	45	66	77	82
32	Shrub and Brush Rangeland	44	64	77	82
33	Mixed Rangeland	46	66	77	83
41	Deciduous Forest Land	31	58	68	75
42	Evergreen Forest Land	35	59	73	79
43	Mixed Forest Land	39	61	74	80
52	Lakes	0	0	0	0
53	Reservoirs	0	0	0	0
61	Forested Wetlands	44	58	68	75
62	Non-forested Wetlands	32	55	68	75
74	Bare Exposed Rock	98	98	98	98
75	Strip Mines	71	80	85	88
76	Transitional Areas	69	78	84	88
81	Shrub and Shrub Tundra	60	74	83	87
82	Herbaceous Tundra	66	76	83	87
83	Bare Ground	74	83	87	90
85	Mixed Tundra	50	65	74	80

## Appendix D. USGS Regression Equations by Region

Table D-1: Region 1 Regression Equations (USGS 1999)

Region 1			
Region Equation	Average Standard Error of Prediction (%)	Equivalent years of record	
$Q_2 = 0.124AREA^{0.845}PREC^{1.44}$	59	0.16	
$Q_5 = 0.629AREA^{0.807}PREC^{1.12}$	52	0.62	
Q <sub>10</sub> = 1.43AREA <sup>0.786</sup> PREC <sup>0.958</sup>	48	1.34	
$Q_{25} = 3.08AREA^{0.768}PREC^{0.811}$	46	2.50	
$Q_{50} = 4.75 AREA^{0.758} PREC^{0.732}$	46	3.37	
$Q_{100} = 6.78AREA^{0.750}PREC^{0.668}$	46	4.19	

Table D-2: Region 3 Regression Equations (USGS 1999)

Region 3			
Region Equation	Average Standard Error of Prediction (%)	Equivalent years of record	
$Q_2 = 0.444AREA^{0.649}PREC^{1.15}$	86	0.29	
$Q_5 = 1.21AREA^{0.639}PREC^{0.995}$	83	.49	
Q <sub>10</sub> = 1.99AREA <sup>0.633</sup> PREC <sup>0.924</sup>	80	.77	
$Q_{25} = 3.37AREA^{0.627}PREC^{0.849}$	78	1.23	
$Q_{50} = 4.70AREA^{0.625}PREC^{0.802}$	77	1.57	
$Q_{100} = 6.42 AREA^{0.621} PREC^{0.757}$	78	1.92	

Table D-3: Region 4 Regression Equations (USGS 1999)

Region 4				
Region Equation	Average Standard Error of Prediction (%)	Equivalent years of record		
$Q_2 = 0.0405AREA^{0.701}(ELEV/1,000)^{2.91}$	64	0.39		
$Q_5 = 0.408AREA^{0.683}(ELEV/1,000)^{2.05}$	57	.95		
$Q_{10} = 1.26AREA^{0.674}(ELEV/1,000)^{1.64}$	53	1.76		
$Q_{25} = 3.74AREA^{0.667}(ELEV/1,000)^{1.24}$	51	3.02		
$Q_{50} = 7.04AREA^{0.664}(ELEV/1,000)^{1.02}$	52	3.89		
$Q_{100} = 11.8AREA^{0.662}(ELEV/1,000)^{0.835}$	53	4.65		

Table D-4: Region 6 Regression Equations (USGS 1999)

Region 6			
Region Equation	Standard Error of Regression (Log Units)	Equivalent years of record	
$Q_2 = 0$			
$Q_5 = 32AREA^{0.80}(ELEV/1,000)^{-0.66}$	1.47	0.233	
$Q_{10} = 590AREA^{0.62}(ELEV/1,000)^{-1.6}$	1.12	0.748	
$Q_{25} = 3,200AREA^{0.62}(ELEV/1,000)^{-2.1}$	0.796	2.52	
$Q_{50} = 5,300AREA^{0.64}(ELEV/1,000)^{-2.1}$	1.1	1.75	
Q <sub>100</sub> = 20,000AREA <sup>0.51</sup> (ELEV/1,000) <sup>-2.3</sup>	1.84	0.794	

**Table D-5: Region 7 Regression Equations (USGS 1999)** 

Region 7				
Region Equation	Average Standard Error of Prediction (%)	Equivalent years of record		
$Q_2 = 0.0150AREA^{0.697}(ELEV/1,000)^{3.16}$	56	0.25		
$Q_5 = 0.306AREA^{0.590}(ELEV/1,000)^{2.22}$	45	1.56		
$Q_{10} = 1.25AREA^{0.526}(ELEV/1,000)^{1.83}$	45	3.07		
Q <sub>25</sub> = 122AREA <sup>0.440</sup>	49	4.60		
Q <sub>50</sub> = 183AREA <sup>0.390</sup>	53	5.27		
Q <sub>100</sub> = 264AREA <sup>0.344</sup>	59	5.68		

Table D-6: Region 8 Regression Equations (USGS 1999)

	Region 8	
Region Equation	Average Standard Error of Prediction (%)	Equivalent years of record
$Q_2 = 598AREA^{0.501}(ELEV/1,000)^{-1.02}$	72	0.37
$Q_5 = 2,620 AREA^{0.449} (ELEV/1,000)^{-1.28}$	62	1.35
Q <sub>10</sub> = 5,310AREA <sup>0.425</sup> (ELEV/1,000) <sup>-1.40</sup>	57	2.88
Q <sub>25</sub> = 10,500AREA <sup>0.403</sup> (ELEV/1,000) <sup>-1.49</sup>	54	5.45
Q <sub>50</sub> = 16,000AREA <sup>0.390</sup> (ELEV/1,000) <sup>-1.54</sup>	53	7.45
$Q_{100} = 23,300AREA^{0.377}(ELEV/1,000)^{-1.59}$	53	9.28

Table D-7: Region 9 Regression Equations (USGS 1999)

Region 9				
Region Equation	Average Standard Error of Prediction (%)	Equivalent years of record		
$Q_2 = 0.0204AREA^{0.606}(ELEV/1,000)^{3.5}$	68	0.14		
$Q_5 = 0.181AREA^{0.515}(ELEV/1,000)^{2.9}$	55	.77		
$Q_{10} = 1.18AREA^{0.488}(ELEV/1,000)^{2.2}$	52	1.70		
$Q_{25} = 18.2 AREA^{0.465} (ELEV/1,000)^{1.1}$	53	2.81		
$Q_{50} = 248AREA^{0.449}$	57	3.36		
Q <sub>100</sub> = 292AREA <sup>0.444</sup>	59	3.94		

NOTE: AREA= Basin Area

PREC=Mean Annual Precipitation

ELEV=Mean Basin Elevation Above Sea Level

## Appendix E. Centerville Creek and Coal Creek Data

## Centerville Creek Data

Table E-1: Centerville Creek Data

Precipitation Gage:	Bountiful Val Verda, Davis County, UT	
Location (lat/long: d,m,s):	40 51 0 111 53 0	
Stream Gage:	Centerville Creek near Centerville, UT	
Location (lat/long: d,m,s):	40 54 59 111 51 44	
Date of Storm:	9/15/2002	
Basin Area (mi^2):	3.17	
Total Precipitation (in):	1.1	
Base Flow (cfs):	1	
Peak Flow (cfs):	2	
Runoff Volume (ft^3):	63270	
UDOT CN:	69.3	
Gaged CN:	69	

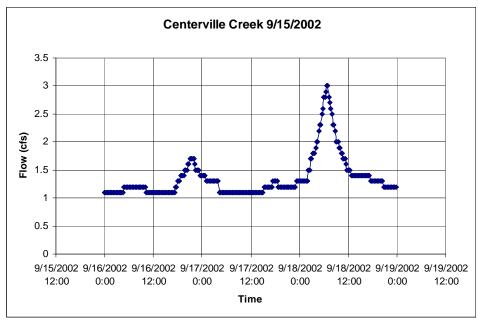


Figure E-1: Centerville Creek Hydrograph

Distance between stream and precipitation gauge: 5 miles.

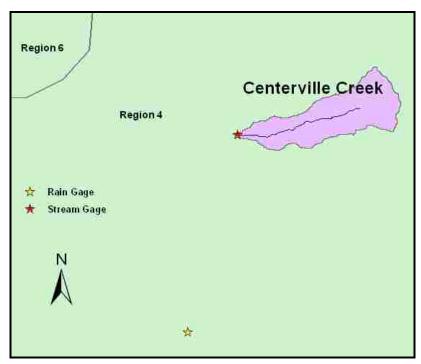


Figure E-2: Centerville Creek Gauge Map

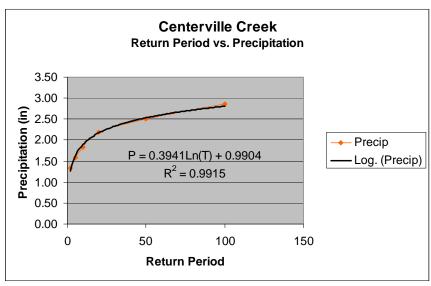


Figure E-3: Centerville Creek Equivalent Return-Period Graph

P = 1.1 inches so the equivalent return-period, T=1.3 years

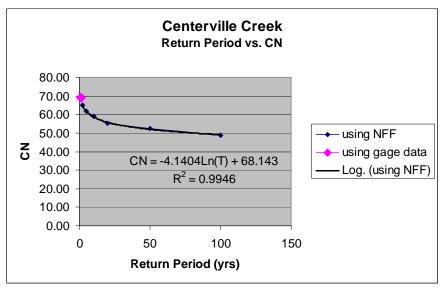


Figure E-4: Centerville Creek Equivalent CN Graph

T=1.3 so USGS regression equation calibrated CN=67.

Gauged CN=69.

2.9% decrease

## Coal Creek Data

Table E-2: Coal Creek Data

Precipitation Gage:	Cedar City 5 E, Iron County, UT	
Location (lat/long: d,m,s):	37 39 0 113 0 0	
Stream Gage:	Coal Creek near Cedar City, UT	
Location (lat/long: d,m,s):	37 40 20 113 2 2	
Date of Storm:	8/23/1987	
Basin Area (mi^2):	77.77	
Total Precipitation (in):	1.3	
Base Flow (cfs):	14	
Peak Flow (cfs):	148	
Runoff Volume (ft^3):	6076000	
UDOT CN:	61.9	
Gaged CN:	69.1	

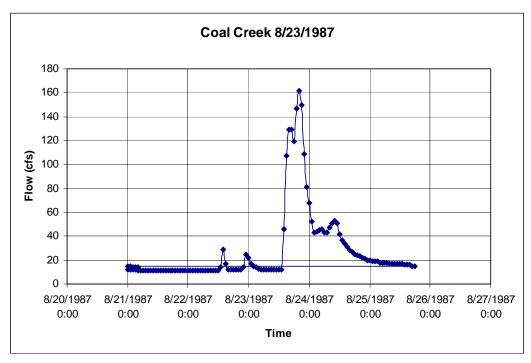


Figure E-5: Coal Creek Hydrograph

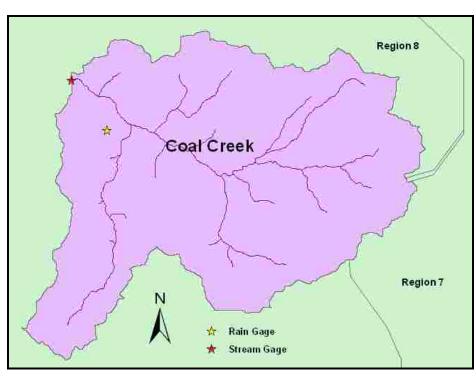


Figure E-6: Coal Creek Gauge Map

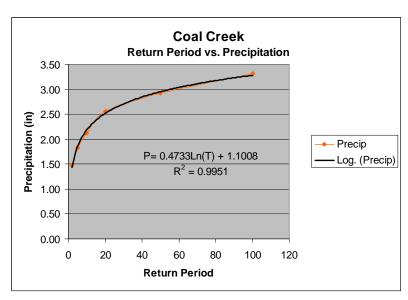


Figure E-7: Coal Creek Equivalent Return-Period Graph

### P = 1.3 inches so the equivalent return-period is T=1.5 years

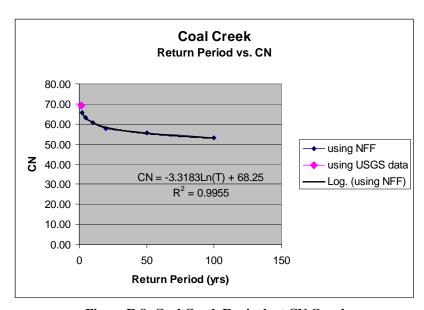


Figure E-8: Coal Creek Equivalent CN Graph

T=1.5 so USGS regression equation calibrated CN=66.9.

Gauged CN=69.1

3.3% decrease

# Appendix F. Guide to Accompanying CD

### **Charlotte Adsero Thesis**

- Thesis
- Thesis Defense Presentation
- Thesis Documents
  - o Utah State Data Comparison
  - o Williams Data Comparison
  - o CN Calibration Table Template
  - o UDOT CN Calibration Results
  - o Utah Watershed Database

### **Associated Research**

- Joel Williams Research
  - o Joel Williams Project Report
  - o Joel Williams Project Data
- Shane Dyer Research
  - o Shane Dyer Project Report
  - o Shane Dyer Project Data